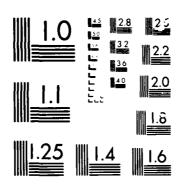
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Monthly and Seasonal Climatology of the Northern Winter over the Global Tropics and Subtropics for the Period 1974 to 1983

Volume IV. 700 mb Winds

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James S. Boyle and C.-P. Chang

May 1986

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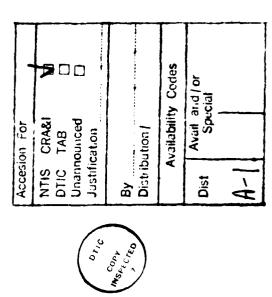
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ABSTRACT

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respective nine-year means are presented for the same variables. The basic wind data used are the analyses, velocity potential and streamfunction from 40°S to 60°N over a global belt for the period 1974 through 1983. In addition the deviations of the individual annual seasonal and monthly means from their This atlas of the 700 mb circulation field contains northern winter monthly and seasonal mean wind operational Global Band Analyses of the United States Navy's Fleet Numerical Oceanography Center.



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1. INTRODUCTION

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Numerical Oceanography Center (FNOC). The procedure used in producing these analyses are described The seasonal calculations are based on the months of December, January, and February. The data used as the basis for these motion fields are the Global Band Analysis (GBA) of the United States Navy's Fleet nine-year seasonal and monthly means and the individual annual seasonal and monthly averages, the This atlas depicts the wintertime (December, January, February) seasonal and monthly mean atmospheric 700 mb motion fields for the period 1974/1975 to 1982/1983. The charts display 700 mb streamfunction, wind vectors, isotachs, and velocity potential from 40°S to 60°N. In addition to the deviations of the individual seasonal and monthly means from their nine-year averages are also presented.

in 1982/83. The latter event is the most intense ENSO event yet observed. Also, the analyses presented here allow the FGGE winter to be placed in a longer term perspective since the FGGE experiment took Krishnamurti et. al. (1983) have presented detailed analysis of the decade prior to 1974. The 1974 - 1983 period contains two 1:1 Nino/Southern Oscillation (ENSO) events, one occurring in 1976/77 and the other The motion fields for the winters from 1974/75 to 1982/83 are of interest since the data cover a period not previously examined in other collections of data. Other works such as Oort (1983) and place in the midst of the period.

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2. DATA SOURCES, ANALYSIS AND COMPUTATIONAL PROCEDURES

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2.1 GLOBAL BAND ANALYSIS

The wind data set used in this work are the operational analyses of the Global Band Analyses of FNOC. These data are produced four times daily by objective procedures on a mercator grid which extends from 40°S to 60°N. The use of the mercator secant projection results in a change in the actual distance between grid points from 140 km at 600N to a maximum value of 280 km at the equator. The objective scheme is designed to take advantage of all the reports in the operational data base, surface synoptic, aircraft, pilot ballons, rawindsonde and satellite data.

interpolate the irregularly spaced data to grid points using a successive corrections technique based on Cressman's (1959) method. The successive corrections method takes several scans through the data by this method. These fields are then adjusted to be consistent with a set of numerical variational analysis adjusted subject to mutual constraints on the fields. However, the surface and 200 mb wind data serve The analysis is performed every six hours for the surface, 700, 400, 250 and 200 mb levels. The first guess field used as input for the objective analysis is the six hour persistence field. The approach is to first reducing the scan radius on each successive scan. Analyses are performed of both wind and temperature (NVA) equations which have incorporated the dynamical constraints of the momentum equations with friction included in the surface layer, (Lewis and Grayson, 1972). Temperature and wind fields are only as a lower and upper boundary condition for the NVA and are not subject to an adjustment. THE STANDARD OF STANDARD STANDARD OF THE STAND

2.2 COMPUTATION OF STREAMFUNCTION AND VELOCITY POTENTIAL

The streamfunction (ψ) and velocity potential (χ) were computed from the following equations:

$$\Delta^2 \chi = \delta$$

$$\nabla^2 \psi = \zeta$$

....

 ζ is the relative vorticity = $\partial v/\partial y \cdot \partial u/\partial x$ and

 δ is the divergence = $\partial u/\partial x + \partial v/c$,. Both ζ and δ were computed using centered differences on the GBA mercator grid taking the appropriate map scale factor into account.

Equation (2) was solved using the boundary condition that $\chi = 0$ on the north and south boundaries which are at 600N and 400S respectively. The method used to compute Ψ was essentially the method II of Shukla and Saha (1974). This technique uses the previously computed values of X to formulate boundary conditions for \\$\psi\$ The depiction of the divergence field appears reasonable away from the boundary. Comparison with the global fields produced by the National Meteorological Center (NMC) for the years since the NMC χ between 40°N and 30°S. Thus the \sqrt{ and \chi} fields in the equatorial regions are sufficently removed from fields have become available indicate that the effect of the boundary conditions on χ is not significant the boundaries that we can assume that these values are not unduly affected by the choice of boundary conditions.

In addition the windfield was directly decomposed into its rotational and divergent components using the method of Endlich (1967). This method does not require any assumption about the boundary conditions. The divergent wind vectors shown in the figures are not computed from X but are those computed by the Endlich technique. The excellent agreement between the χ field and the divergent winds over the entire grid, gives confidence in the accuracy of the computations.

3. DISCUSSION

The winter (DJF) mean nine year data (Figs. A1 - A3) are in reasonable agreement with the data of Newell et. al. (1974) and Oort (1983). The 700 mb wind field field is not a common field in most data collections. The familiar features of all these data compilations are prominent in the present work. The Africa and a zonally elongated maxima centered on New Guinea. This pattern is similar to the surface χ X field (Fig. A3) shows a broad band of equatorial convergence, with maxima over South America,

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There is a rapid decrease in the gradient of χ and thus the divergent wind magnitude in going from the West Pacific. This is in agreement with the findings of Thompson et. al. (1979), who found that the divergence diminished rapidly with height in the east Atlantic but was nearly as strong at 700 mb as at the field shown in Boyle and Chang (1986), but the gradients and divergent wind is much weaker at 700 mb. surface to 700 mb. The largest values of χ and divergent wind at 700 mb are found in the Equatorial South America have their lower level convergence largely restricted to below 700 mb. These features are surface in the west Pacific. It would appear that two major convective centers, equatorial Africa and consistent with the tropical outgoing longwave radiation (OLR) data, Boyle and Lau (1984). However, the South Pacific convergence zone (SPCZ) which is in evidence in the OLR data and the 200 mb χ field does not appear to be a distinct feature in the 700 mb x field. Figure I is a plot of the magnitude of the low level (surface to 700 mb) wind shear. In regions continents. There are also large values across the Northern Africa region. The positioning of the strong where the geostrophic approximation is valid, this is related to the mean low level thermal gradient. Not unexpectedly, the largest values are found along and to the east of the Asian and North American baroclinicity in the western ocean basins is consistent with these regions as being considered as source regions for baroclinic waves. In the Southern Hemisphere (summer) the strongest gradients tend to be on the west coasts of the continents, evidently related to the very cool waters off these coasts. Figure II is a time, longitude plot of X winter, seasonal anomalies along the Equator. The sense of anomalies is that positive anomaly maxima imply convergence relative to the mean. In contrast to the surface and 200 mb, Boyle and Chang (1986), Boyle and Chang (1984), the 700 mb anomalies do not show large deviations during the 1982/83 ENSO event. Evidently, the major circulation changes are above and below 700 mb leaving this level largely unchanged. Examination of the anomalies do not indicate as good a relationship between the χ_{700} and the OLR anomalies as exhibited between the χ_{200} and the OLR anomalies.

The midlatitude flow in Fig. V has the largest variations on the eastern sides of the major ocean basins in 400N, respectively. In Fig. III the 82/83 ENSO event is barely in evidence. As in the χ field, the 700 mb The subtropical flow of Fig. IV indicates that the most active region at this level is over the Pacific basin. the zonal wind field. The meridional wind anomalies have a wavelike nature to them east of about $180^{
m O}$ Further overviews of the interannual, seasonal variations are provided by Figs. III, IV and V. These are time, longitude plots of the 700 mb wind winter season anomalies along the Equator, 200N, and zonal wind (Fig. IIIa) shows almost no sign of the rather dramatic anomalies observed at the surface. to about 30°E. i programa e especial de deserva especial de la companya de la com

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FI GURES

FIGURE 1

Nine winter season (1974/75 - 1982/83) mean wind shear speed for the layer from the surface to 700 mb. Contour interval for the wind shear magnitude is 2 ms⁻¹. The black dots on the plot indicate terrain heights greater or equal to 1500 m, smoothed to the GBA grid.

FIGURE 11

Time versus longitude plot of winter season anomalies of 700 mb velocity potential averaged from 5°N to 5°S about the Equator. The ordinate labels refer to the year of the January of the winter involved. The contour interval is $0.25 \times 10^6 \text{ ms}^{-2}$. Dashed lines are negative, solid lines are positive values. The zero contour is supressed. Medical Indiana Incomes pageses Relations pages process

FIGURE III

(a) As in Fig. I except for the zonal wind component. The contour interval is 0.25 ms⁻¹ (b) As in (a) except for the meridional wind component.

FIGURE IV

As in Fig. II except the reference latitude is 200N.

FIGURE V

As in Fig. II except the reference latitude is 40°N.

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FIGURES AI - A3

Nine winter season (1974/75 - 1982/83) mean circulation fields at the 700 mb. Variables Analysis mercator grid are shown on the left hand side and bottom of each figure. The longitude grid is marked every 300 from the Greenwich meridian on the extreme left and the latitude is marked every 100 from 400S at the bottom of the figure. The black dots on the displayed are wind vectors and isotachs, streanfunction, and velocity potential. Contour interval for the streamfunction is $2.0 \times 10^6 \,\mathrm{m}^2\mathrm{s}^{-1}$, for the velocity potential it is $1.0 \,\mathrm{x}$ 10⁶ m²s⁻¹, and for the isotachs it is 2.5 ms⁻¹. The vector scale is given in the upper right portion of the wind vector isotach plots. The grid intervals of the FNOC Global Band wind vector plots indicate terrain heights greater or equal to 1000 m, smoothed to the GBA grid. The contour plots use the convention that negative values are dashed, positive values are solid. Although the sign of the streamfunction and velocity potential has no meaning in itself, this plotting convention allows the principle maxima and minima in these fields to be more readily discerned. TOCOTOGO POSSOSO PASSOSOS PASSOSOS POSSOSOS PASSOSOS PASSOSOS PASSAS PASSOS PAS

FIGURES BI TO B27

are solid. The zero contour is not drawn. The vector scale is given in the upper right part of contours corresponding to negative values are dashed, those corresponding to positive values potential it is $1.0 \times 10^6 \text{ m}^2\text{s}^{-1}$, and for the isotachs it is 1.5 ms^{-1} . On the deviation plots the 10^6 m²s⁻¹, for velocity potential it is 1.0 x 10^6 m²s⁻¹, and for the isotachs it is 2.5 ms⁻¹. For the deviation fields the contour interval for the streamfunction is 1.0 x $10^6 \, m^2 s^{-1}$, for velocity and velocity potential. For the mean fields the contour interval for the streamfunction is 2.0~xJanuary of the winter. Variables displayed are the wind vectors and isotachs, streamfunction, Figs. Al to A3). The figures are labeled with the year corresponding to the year of the Individual winter season mean and deviation circulation fields at the 700 mb for the winters from 1974/75 to 1982/83. The deviations are differences from the nine-year seasonal mean (

the wind vector plots.

FIGURES CI TO C9

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Nine year (1974 to 1983) monthly mean circulation fields at the 700 mb. The months isotachs, streamfunction, and velocity potential. The contour interval for the streamfunction is displayed are December, January, and February. Variables displayed are the wind vectors and $2.0 \times 10^6 \text{ m}^2\text{s}^{-1}$, for velocity potential it is $1.0 \times 10^6 \text{ m}^2\text{s}^{-1}$, and for the isotachs it is 2.5 ms^{-1} The vector scale is given in the upper right part of the wind vector plots.

FIGURES DI TO D81

Individual monthly mean and deviation circulation fields at the 700 mb for the months from December 1974 to February 1983. The deviations are differences from the nine year monthly mean (Figs. C1 to C9). The months displayed are December, January, and February. potential it is 1.0 x 10^6 m²s⁻¹, and for the isotachs it is 2.5 ms⁻¹ For the deviation fields the For the mean fields the contour interval for the streamfunction is $2.0 \times 10^6 \text{ m}^2\text{s}^{-1}$, for velocity contour interval for the streamfunction is 1.0 x $10^6 \, \mathrm{m}^2 \mathrm{s}^{-1}$, for velocity potential it is 1.0 x 106 m²s⁻¹, and for the isotachs it is 1.5 ms⁻¹. On the deviation plots the contours Variables displayed are the wind vectors and isotachs, streamfunction, and velocity potential. The zero contour is not drawn on the deviations. The vector scale is given in the upper right corresponding to negative values are dashed, those corresponding to positive values are solid. part of the wind vector plots. on source of the second of the second of the second of the second property of the second of the seco

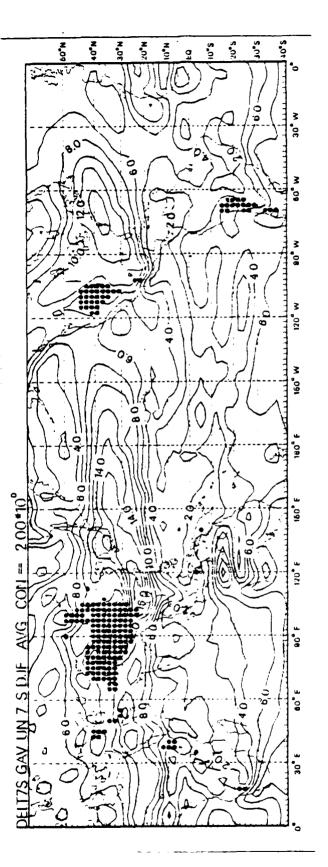
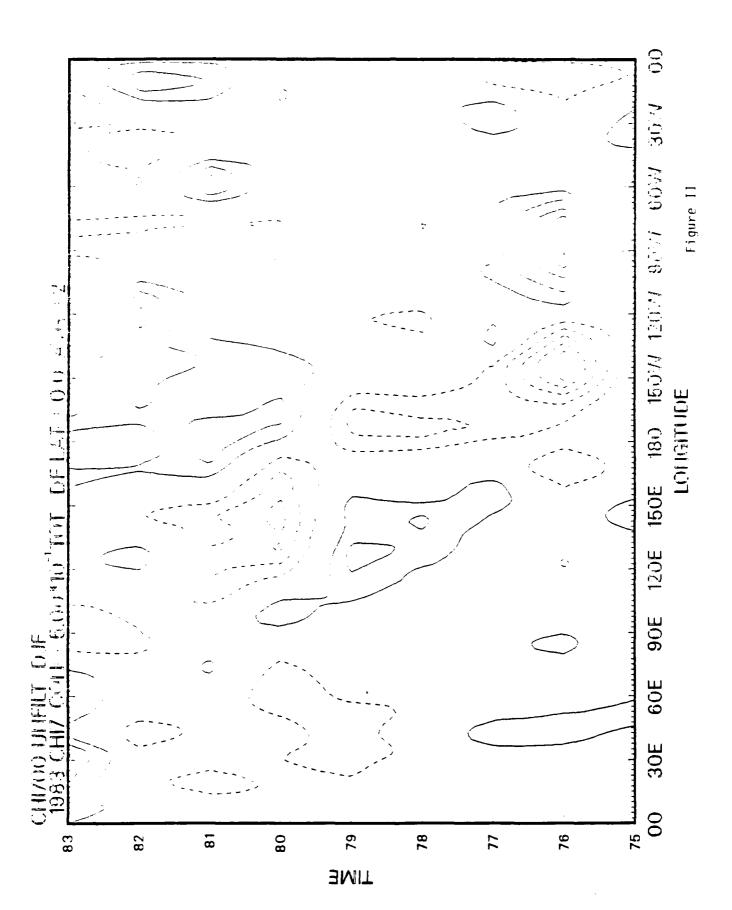
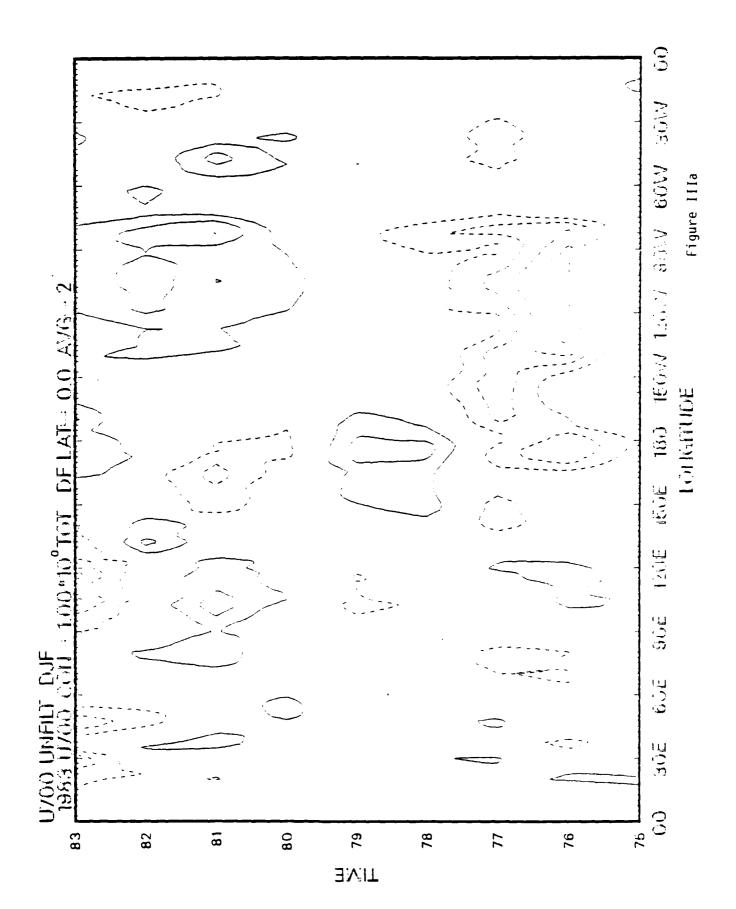
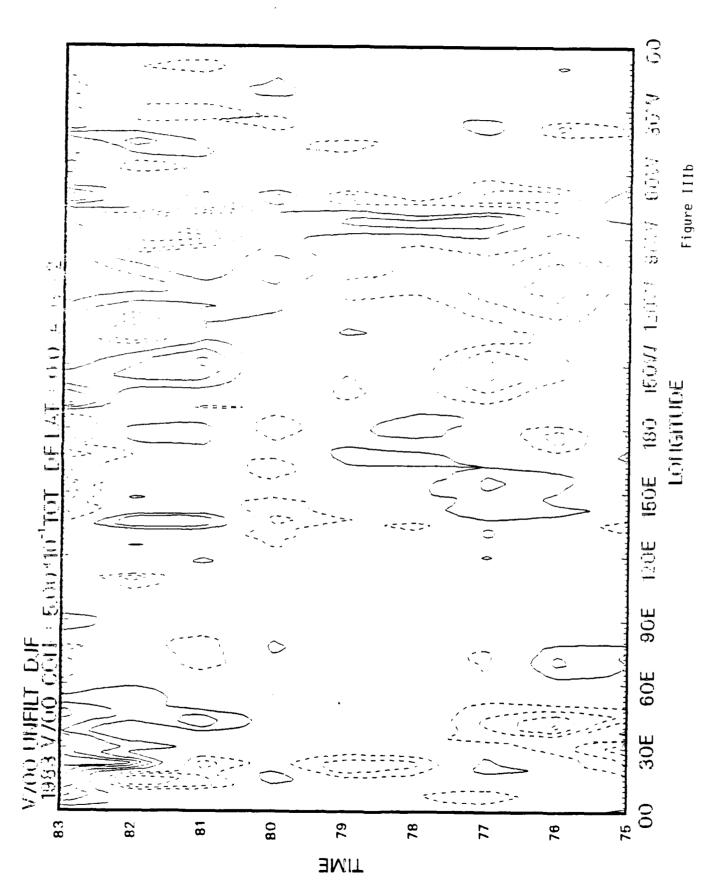
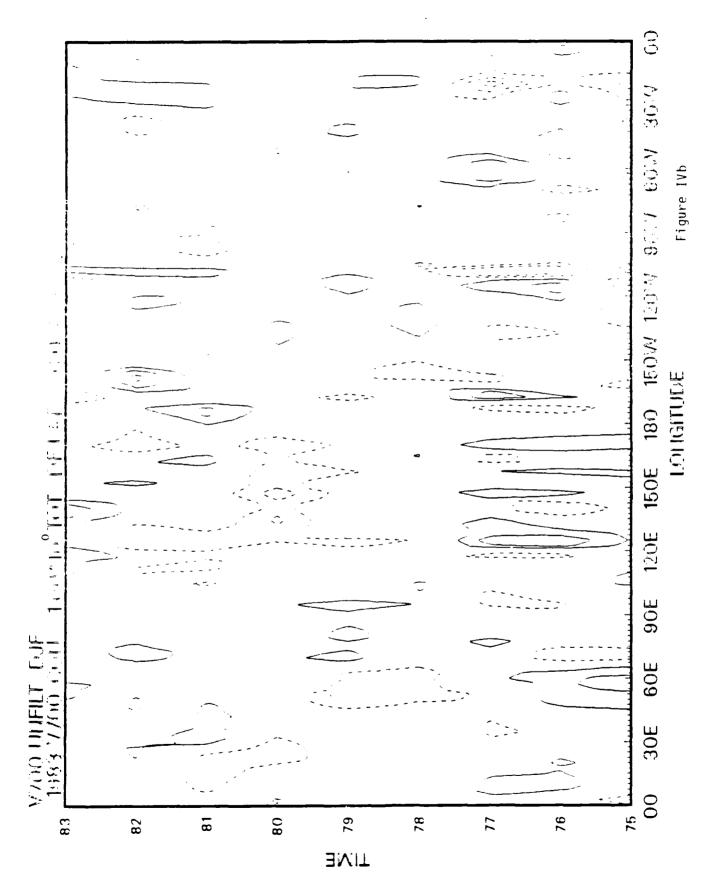


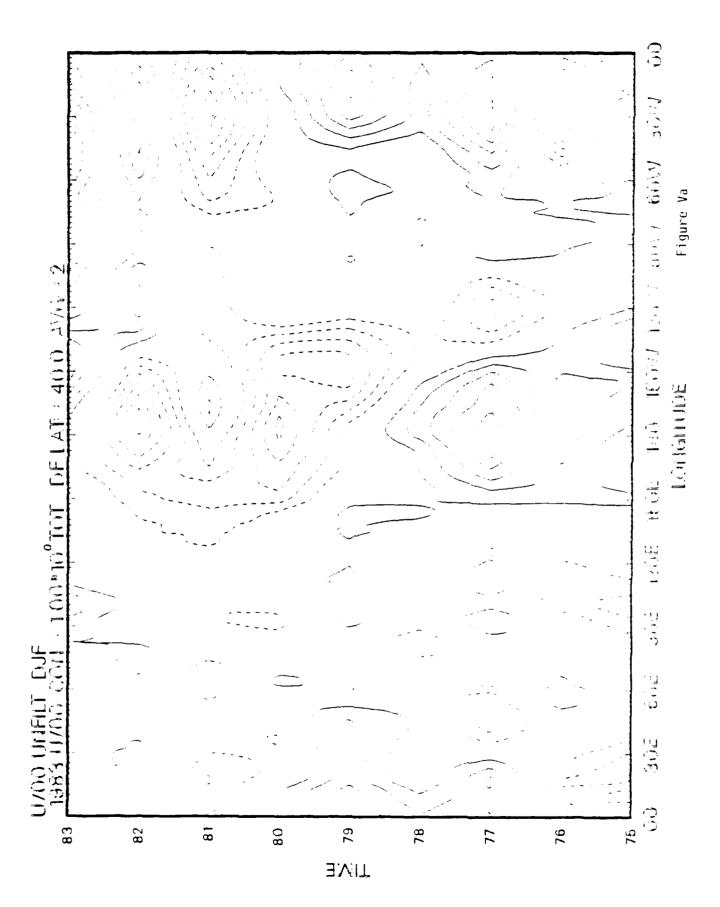
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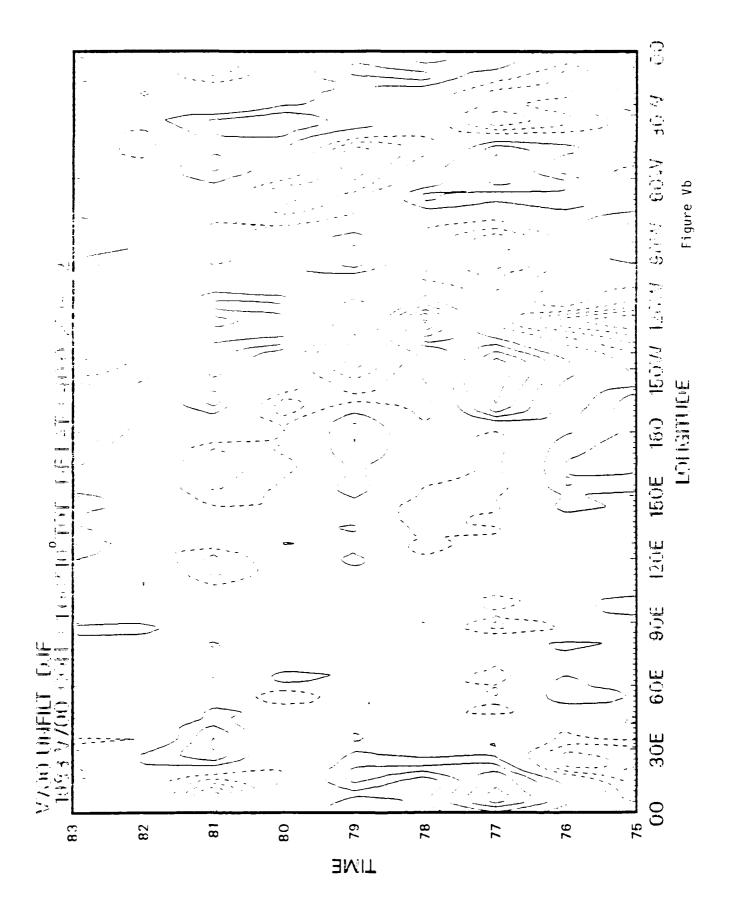


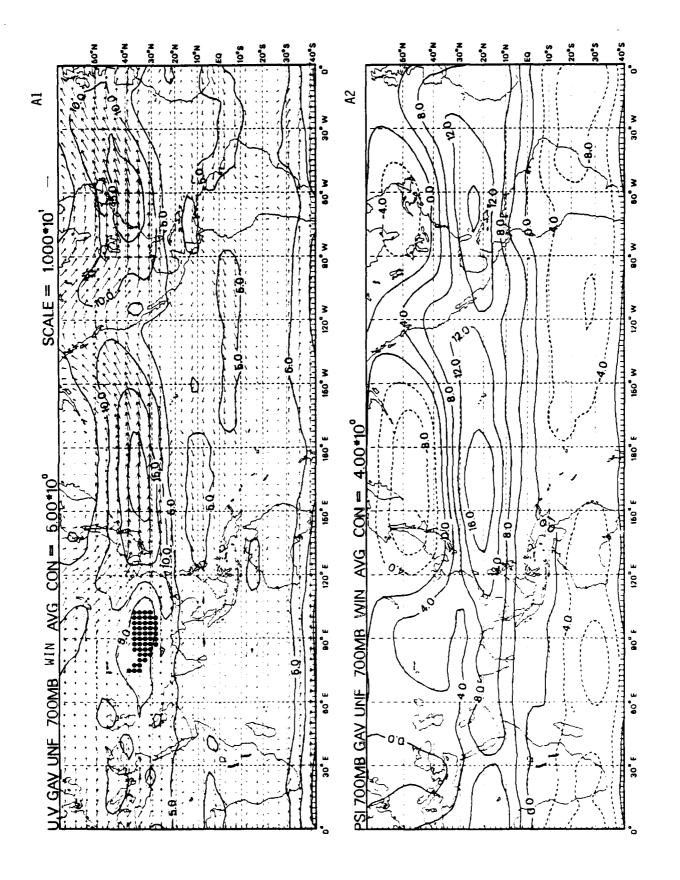


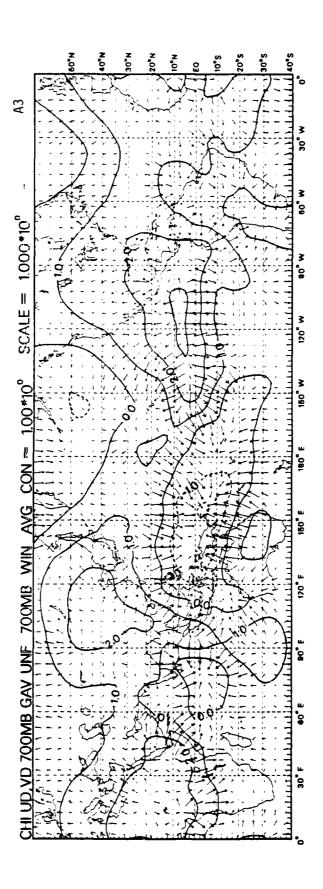


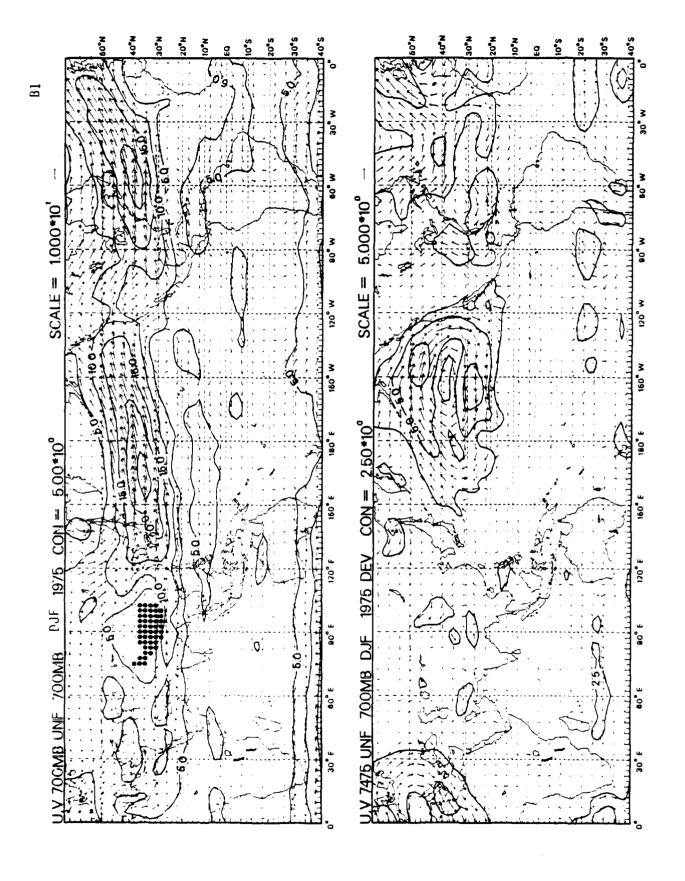


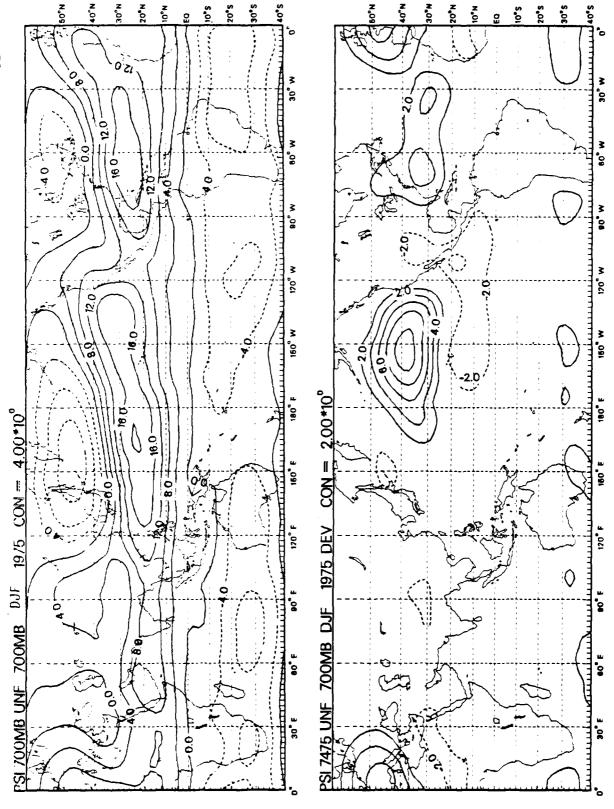


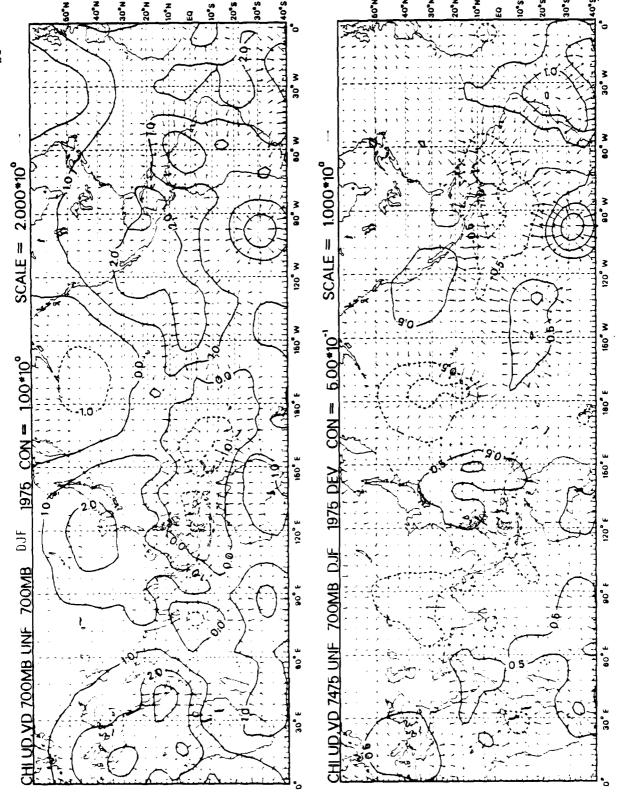


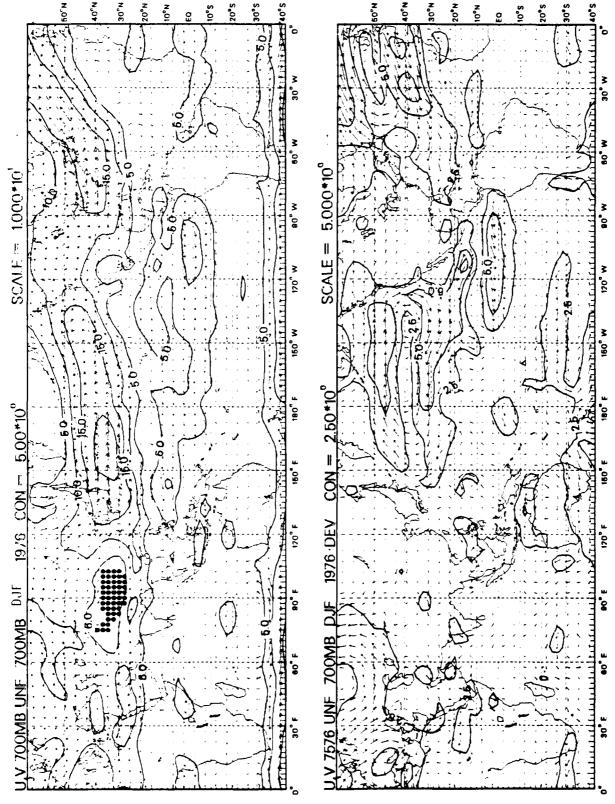


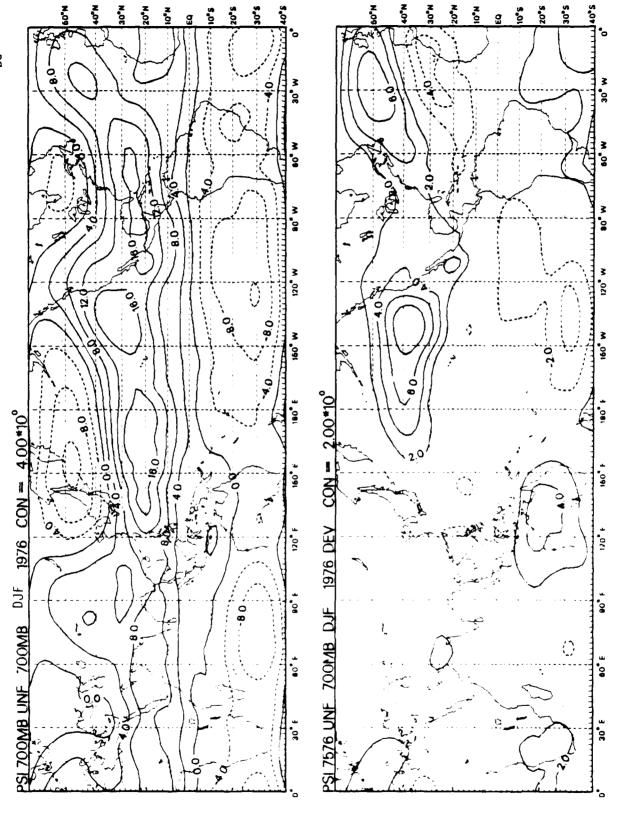


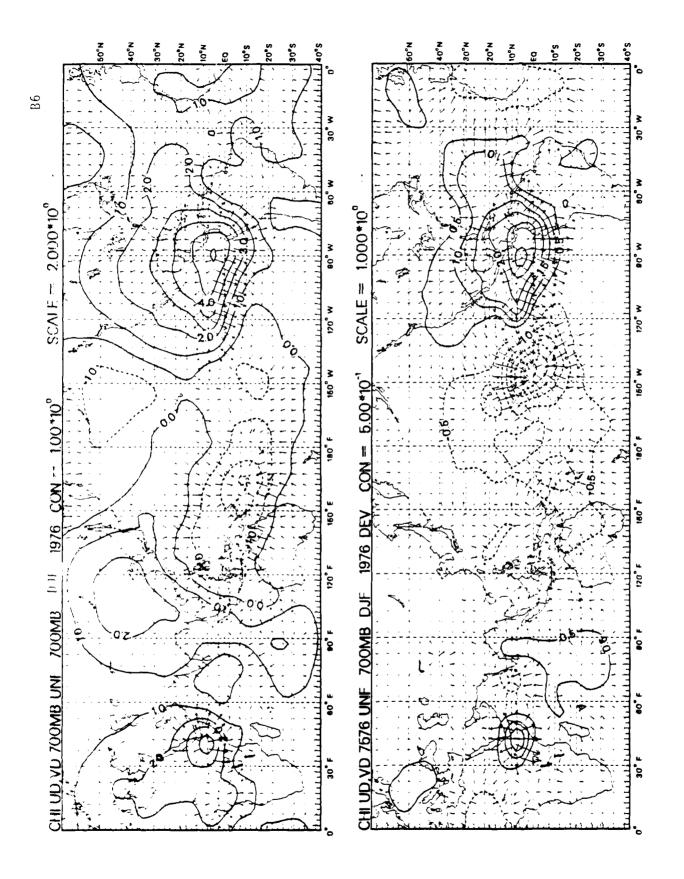


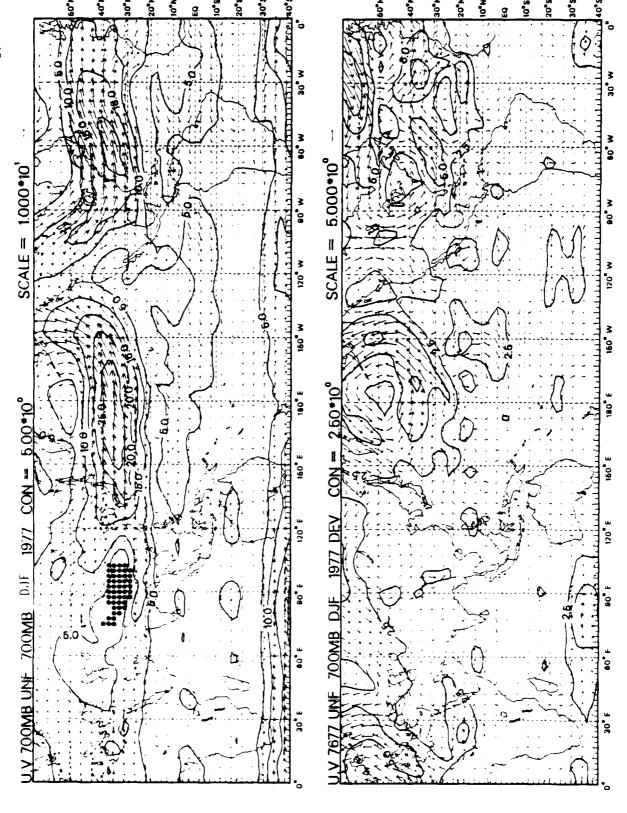


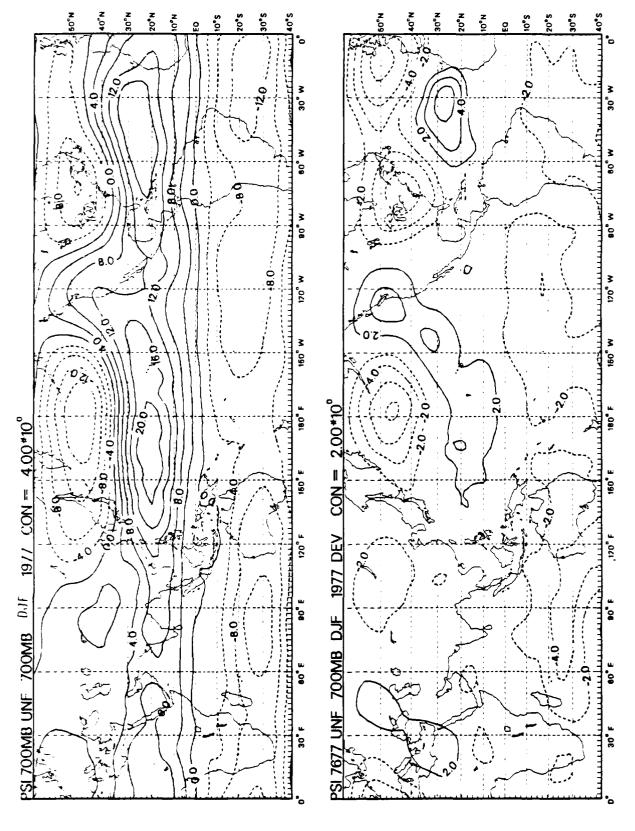


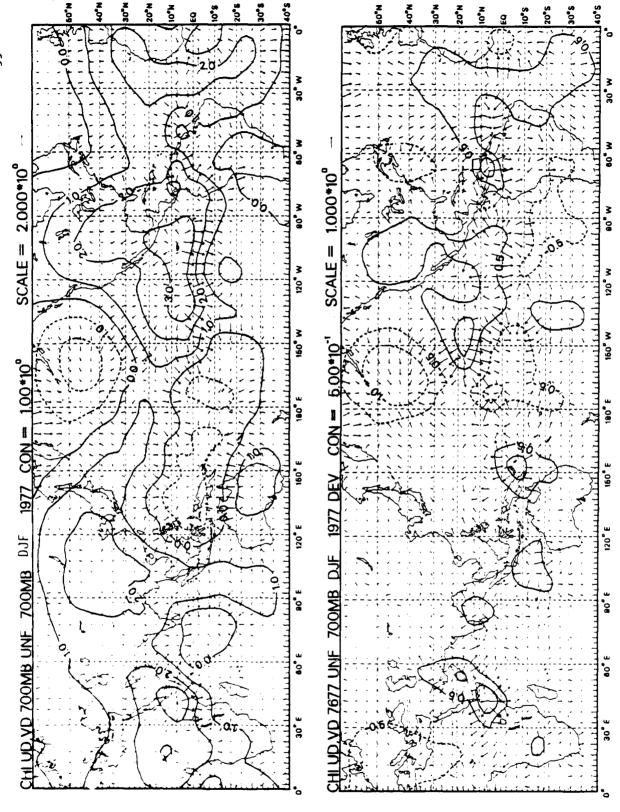


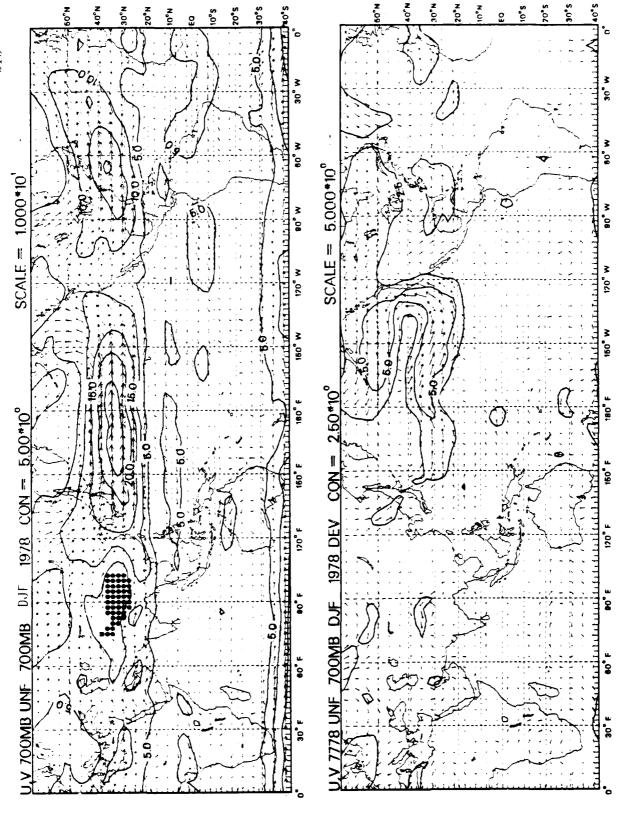


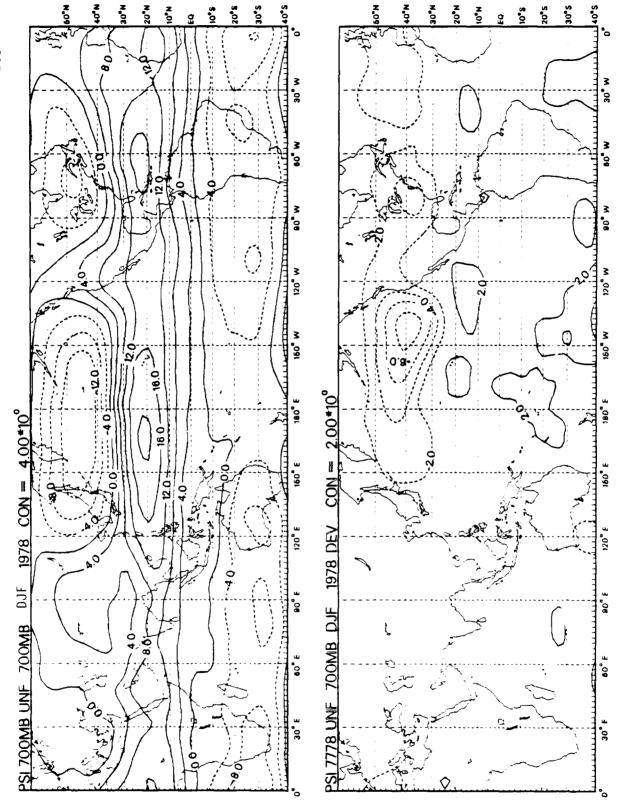


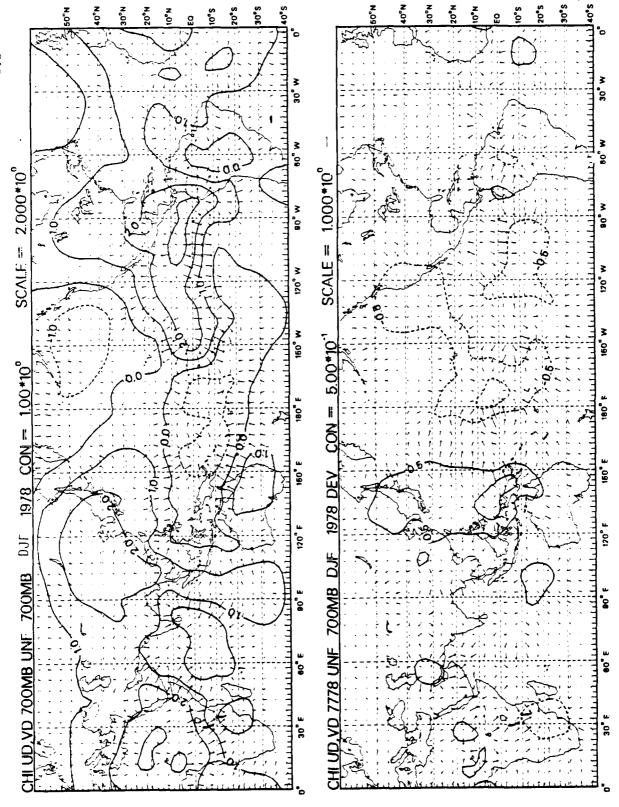




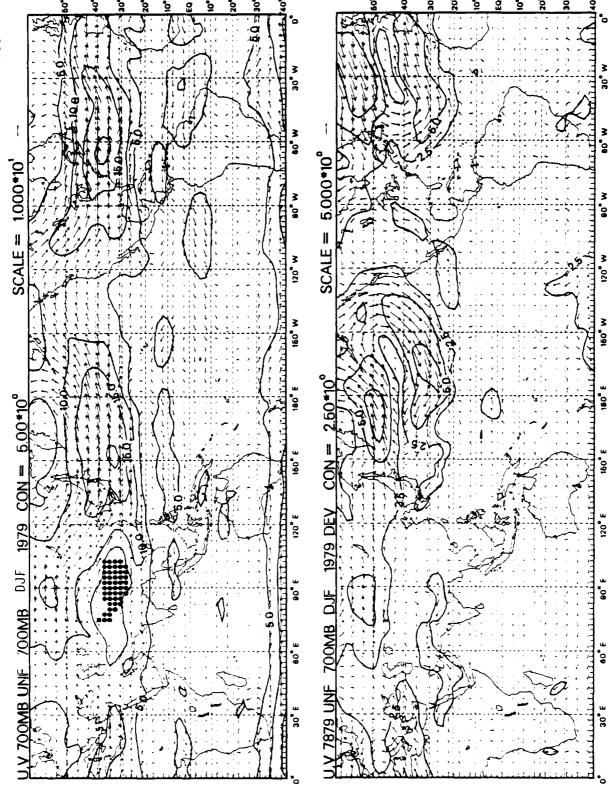


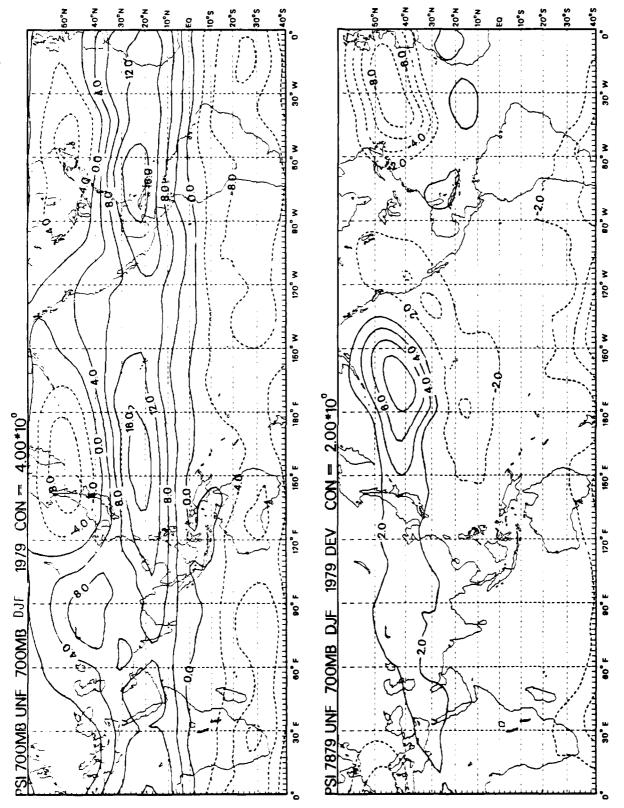


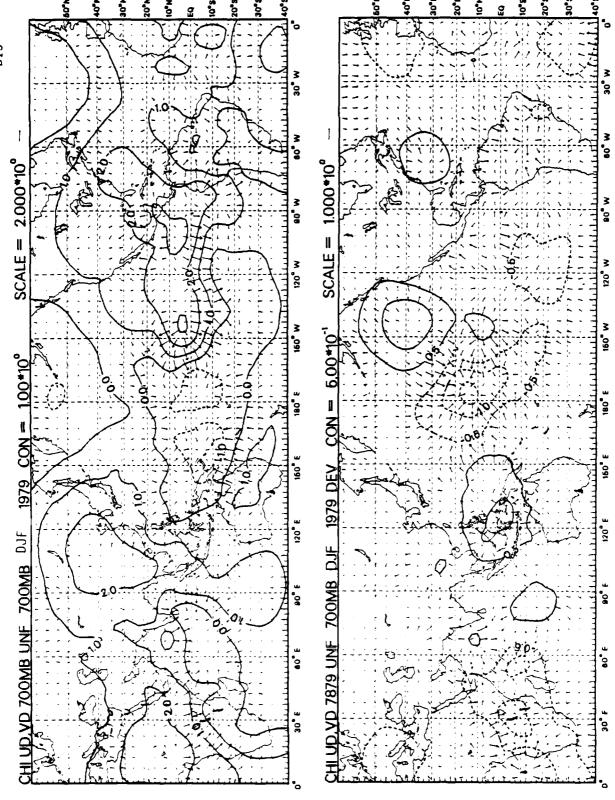


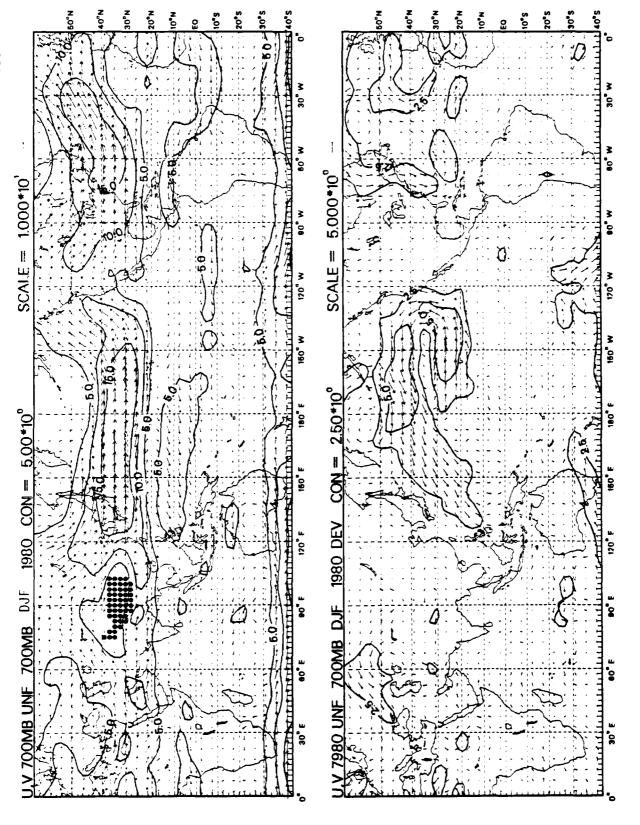


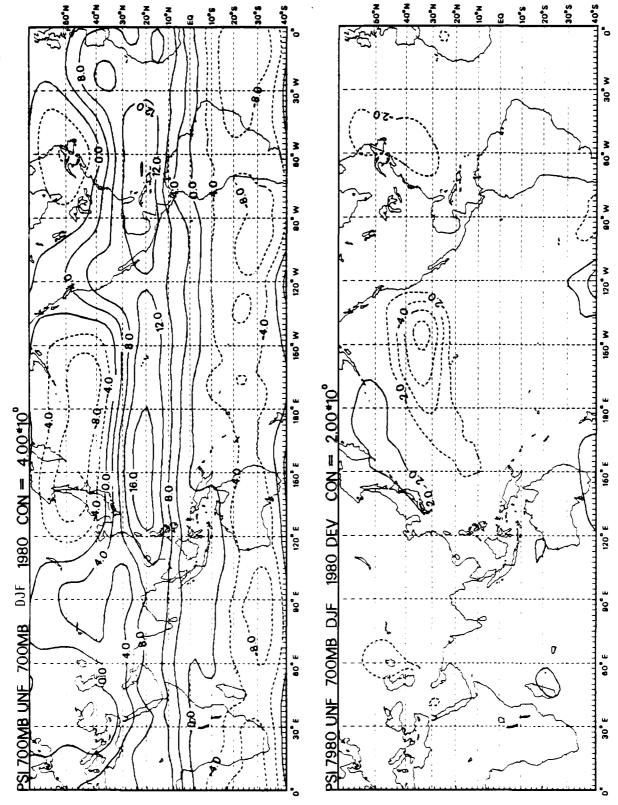


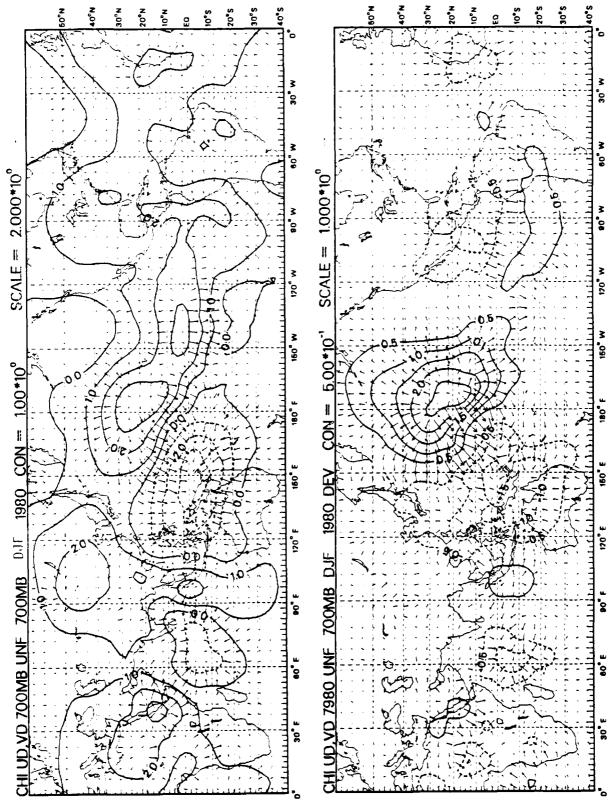


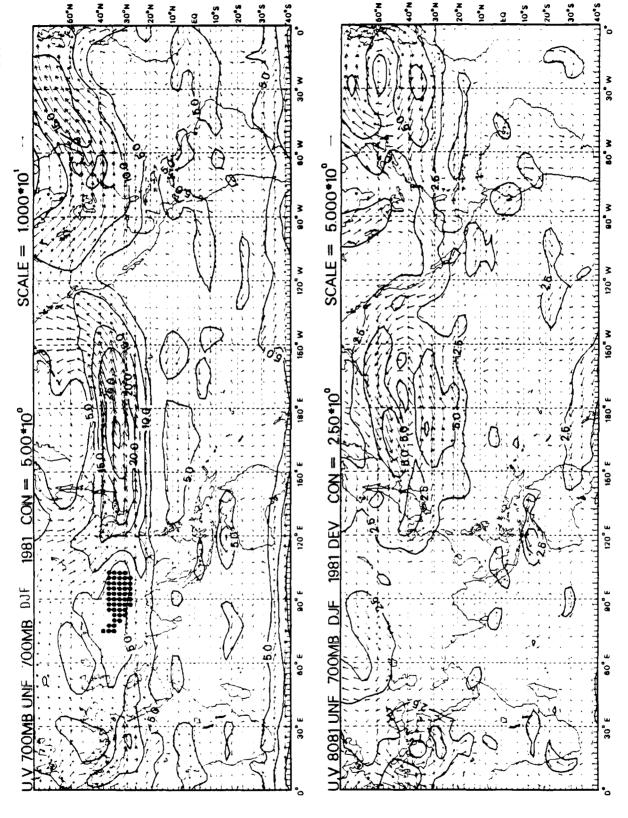


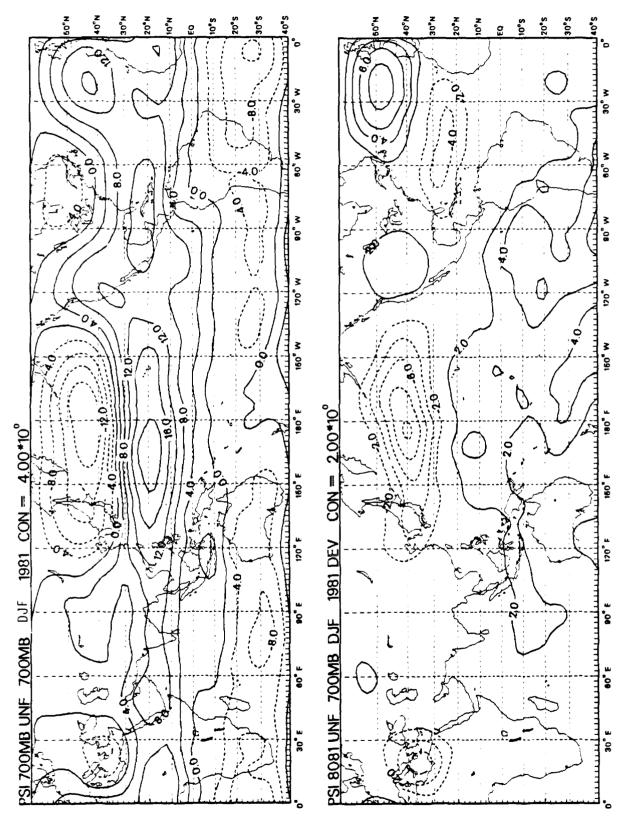




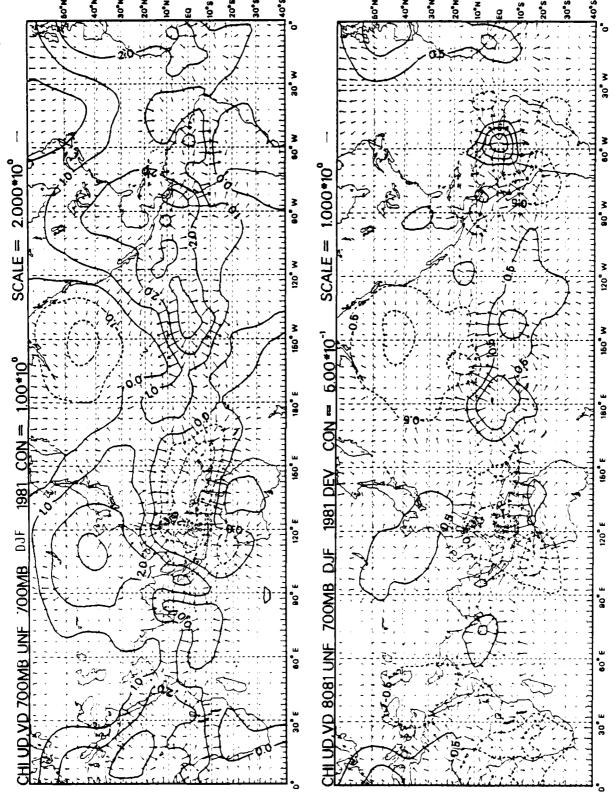


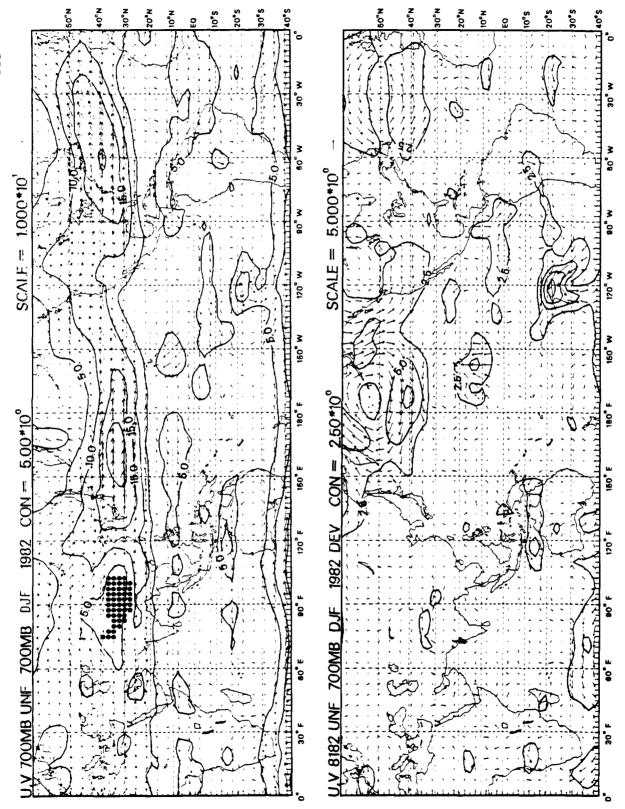


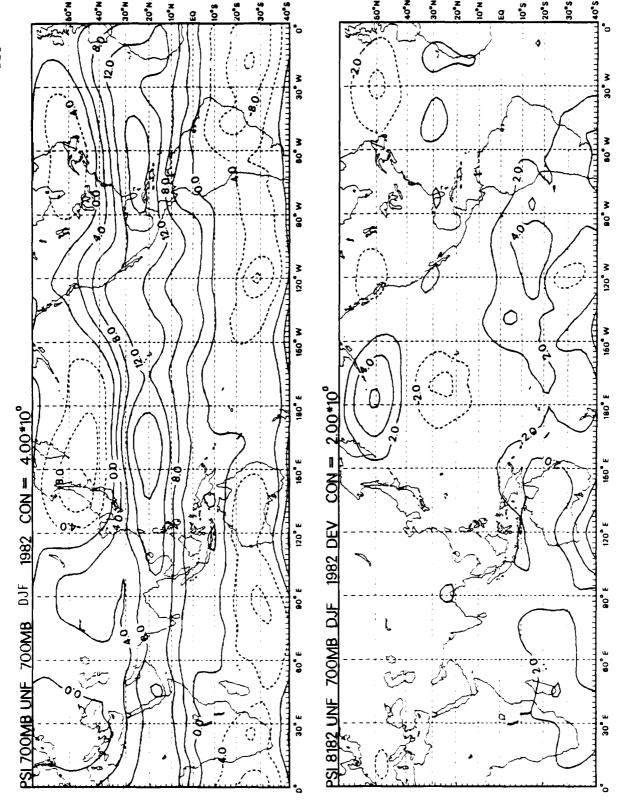


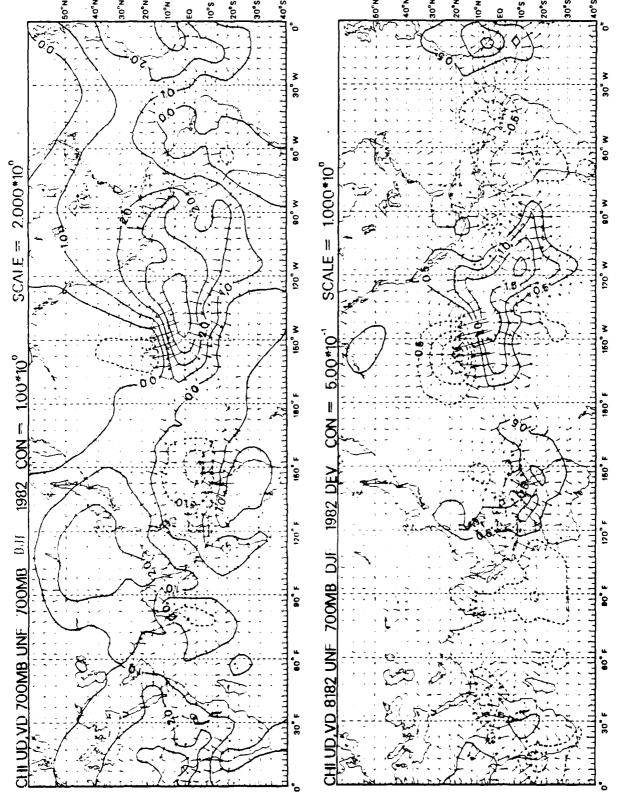


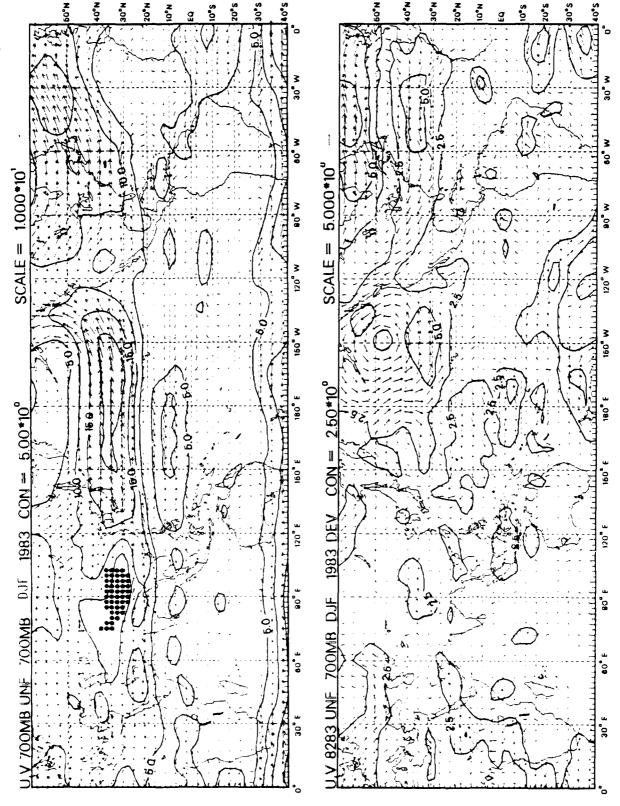


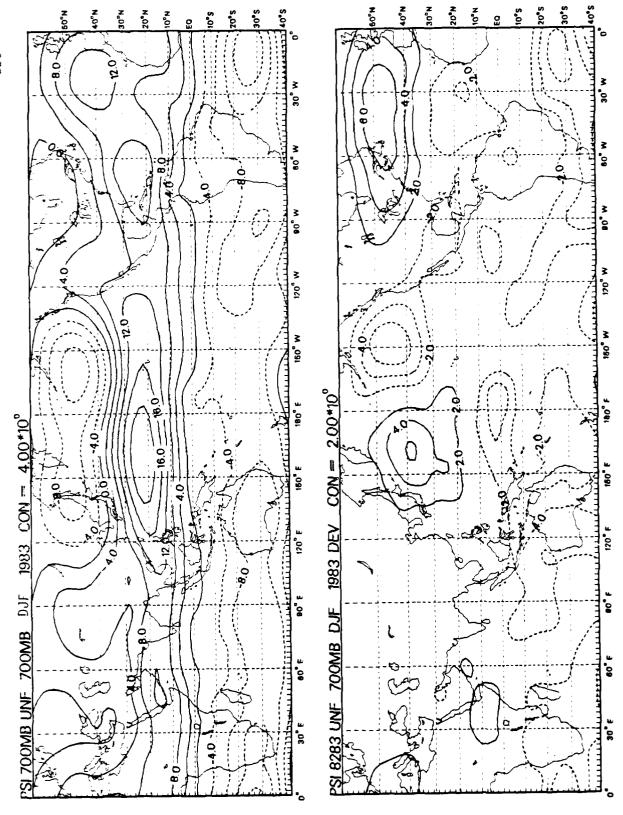






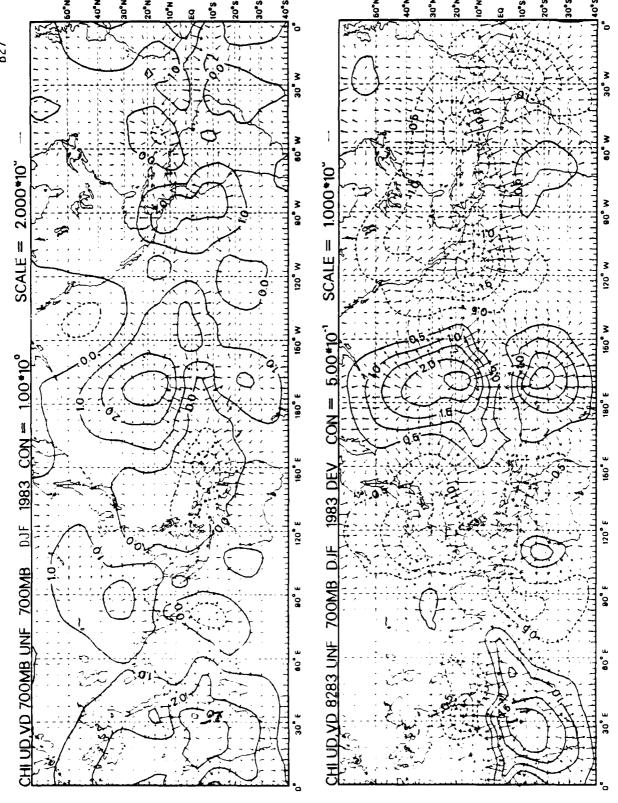


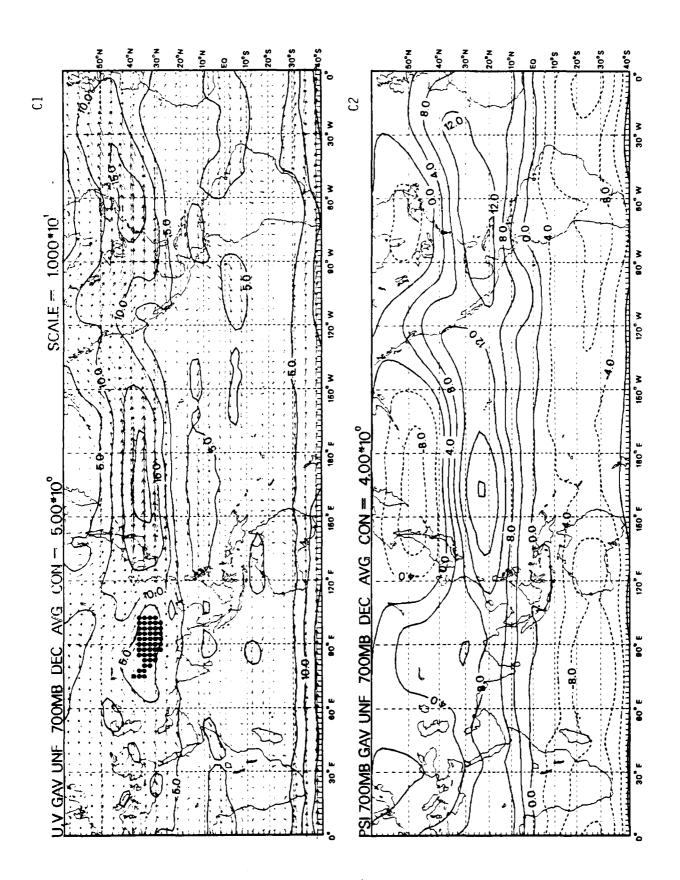


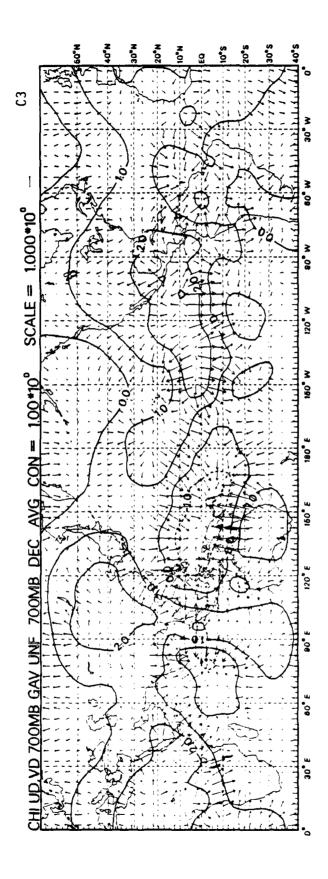


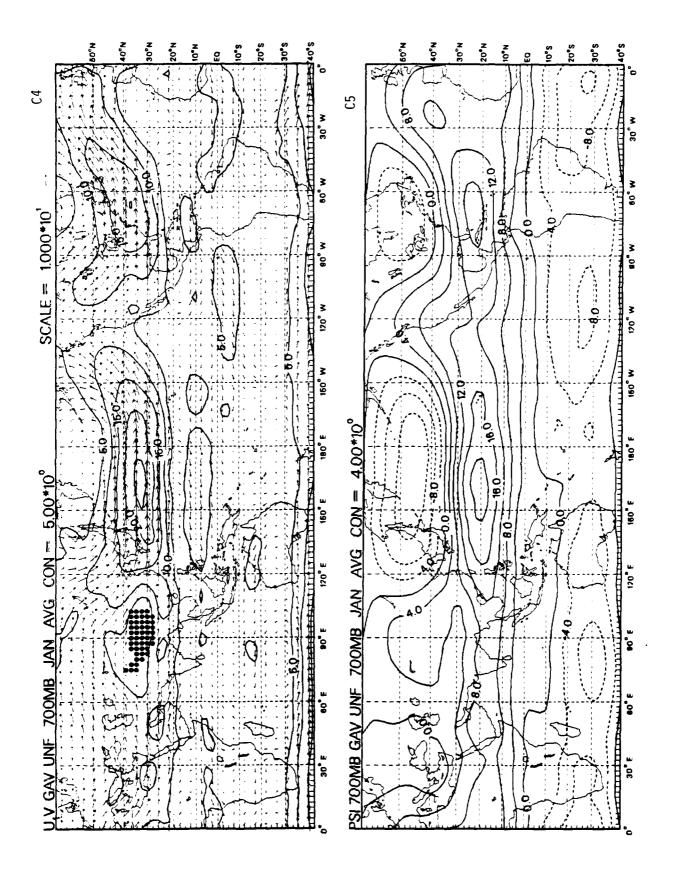
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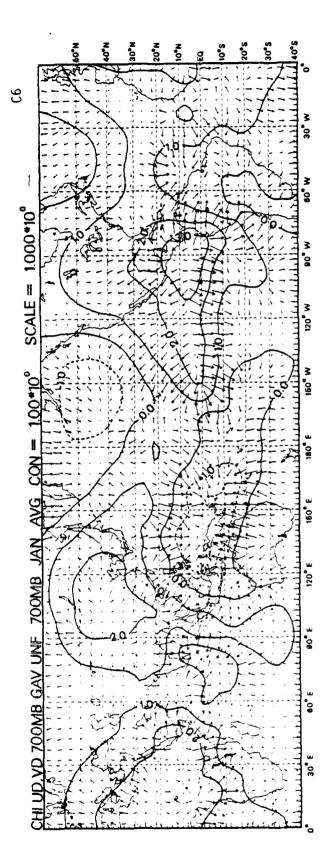












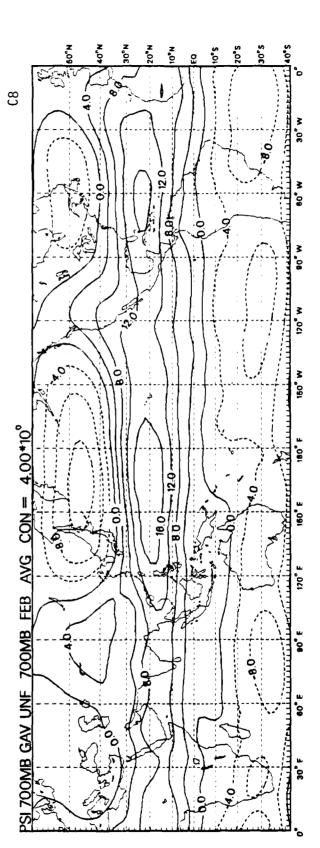
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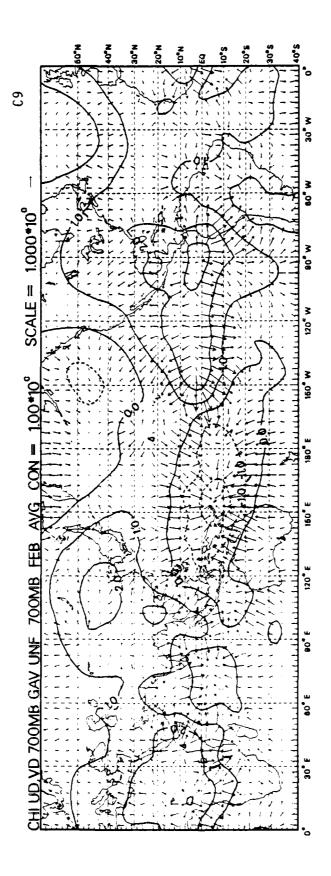
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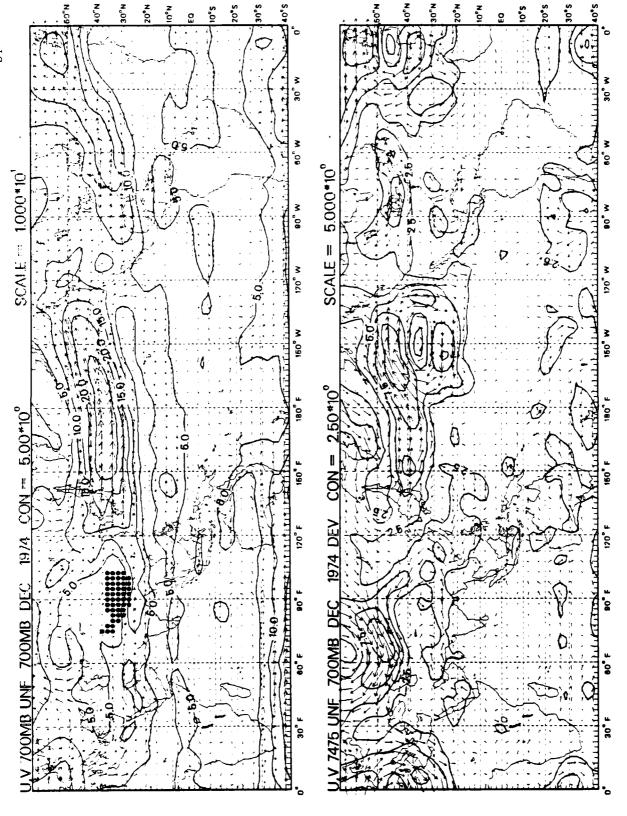
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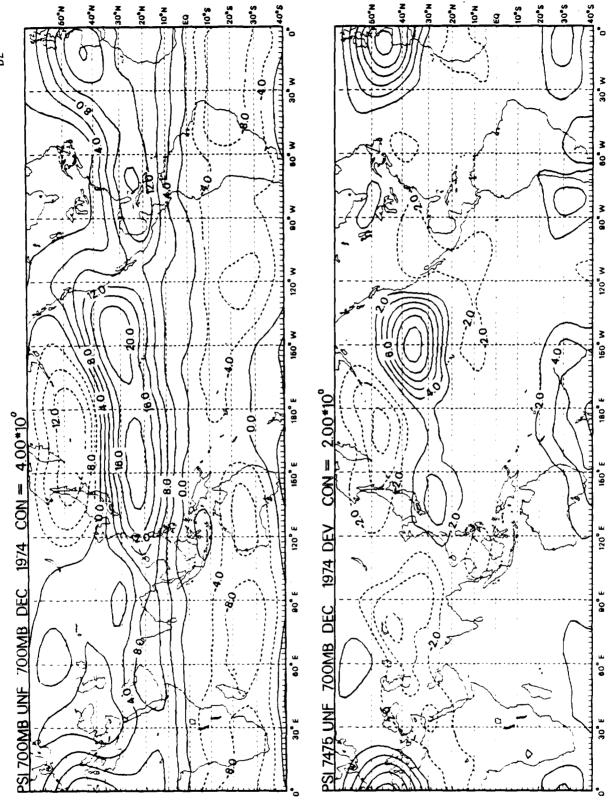
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180° E

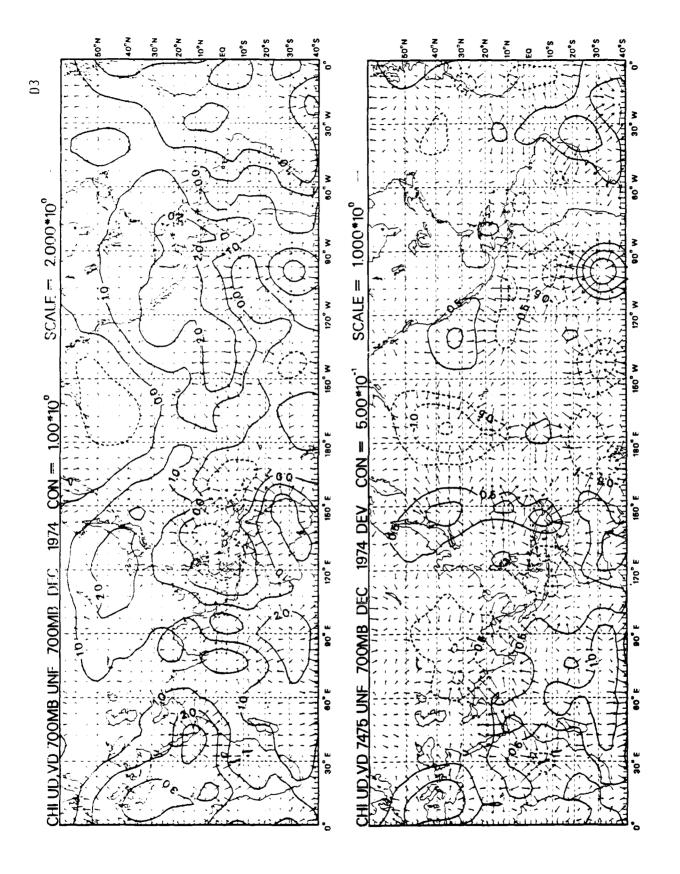


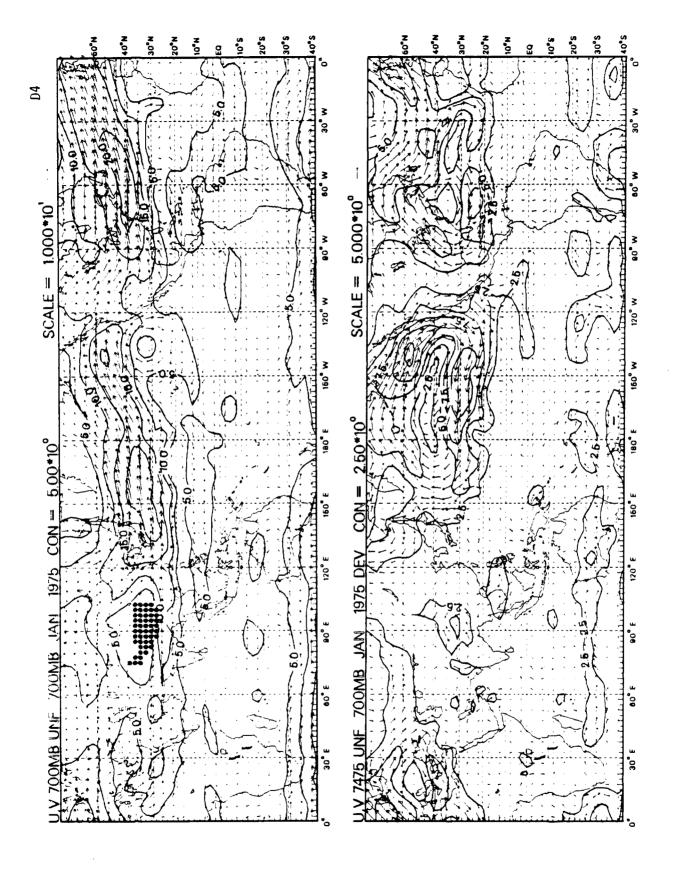


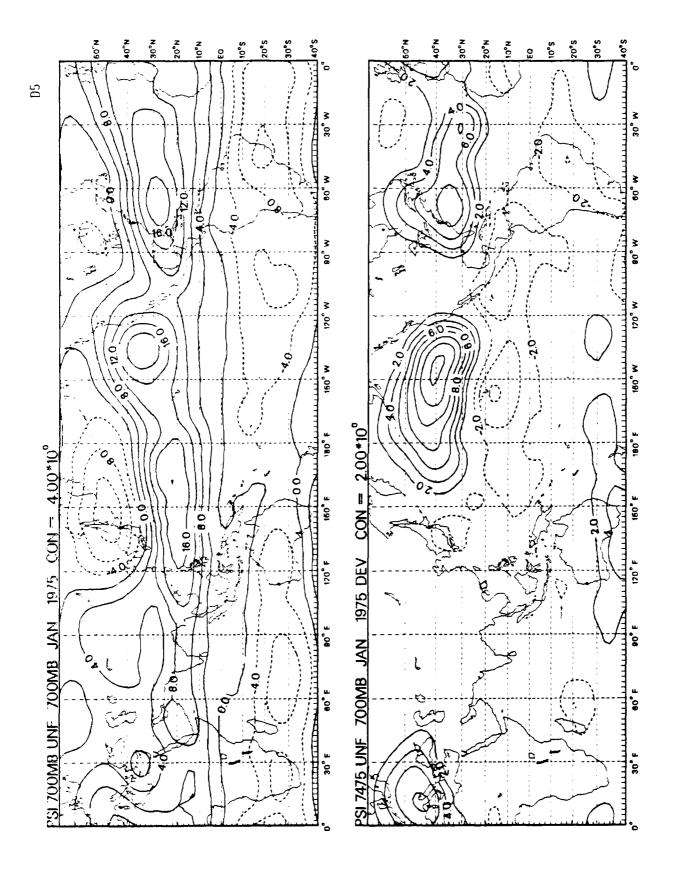


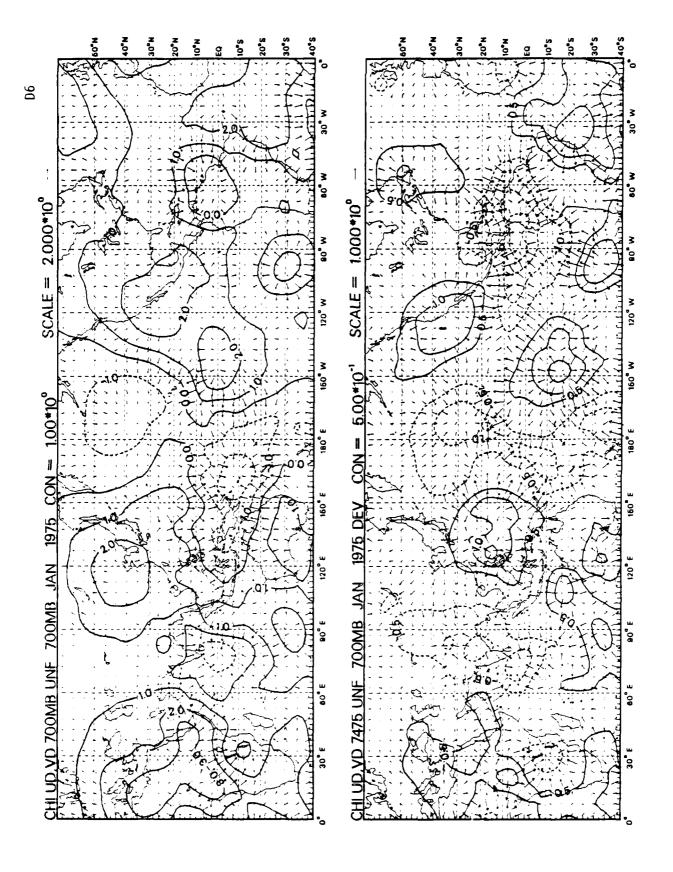


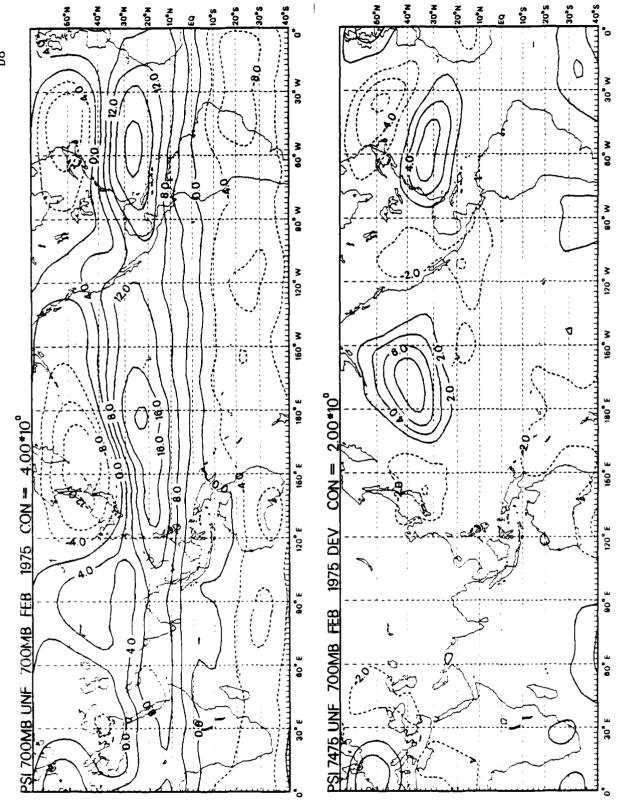
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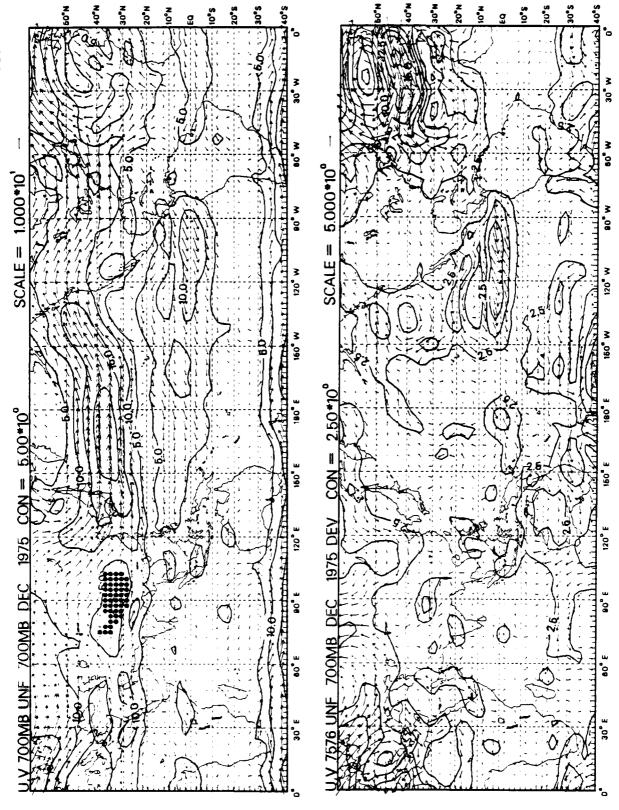




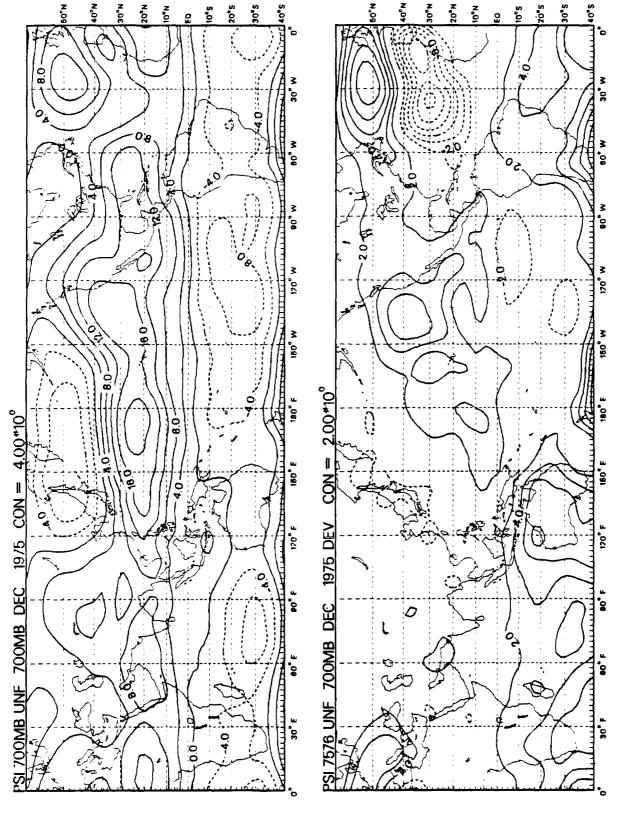




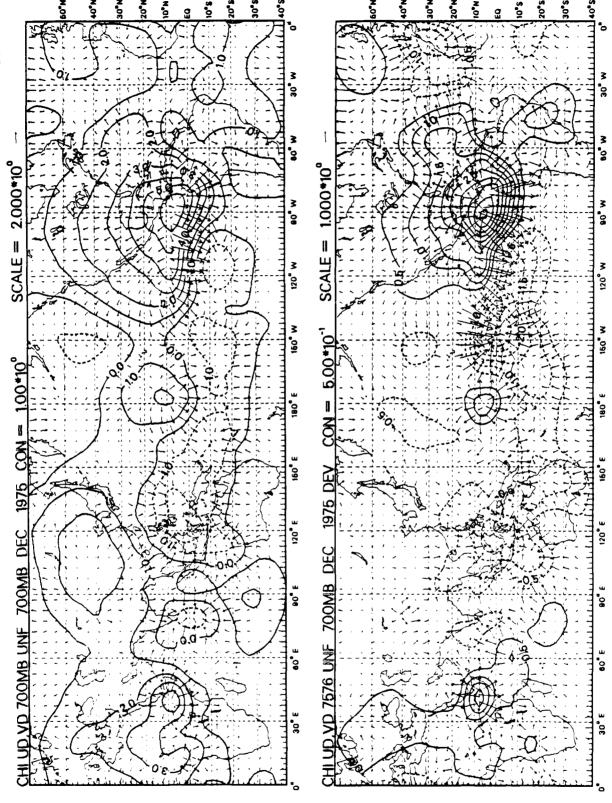


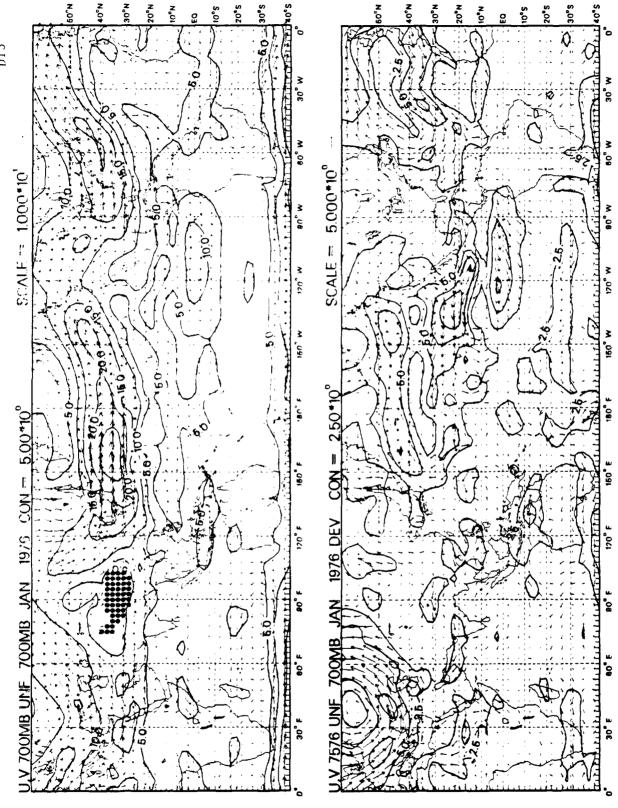


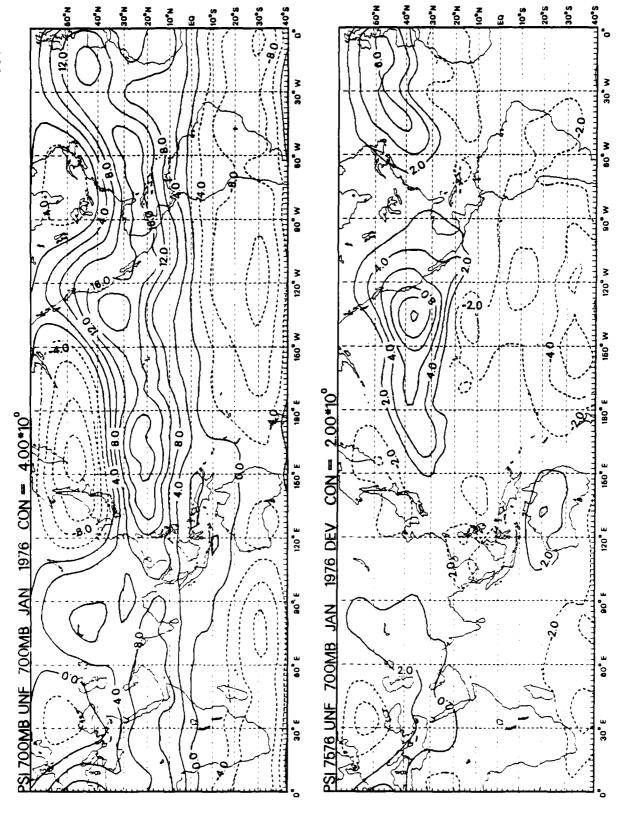
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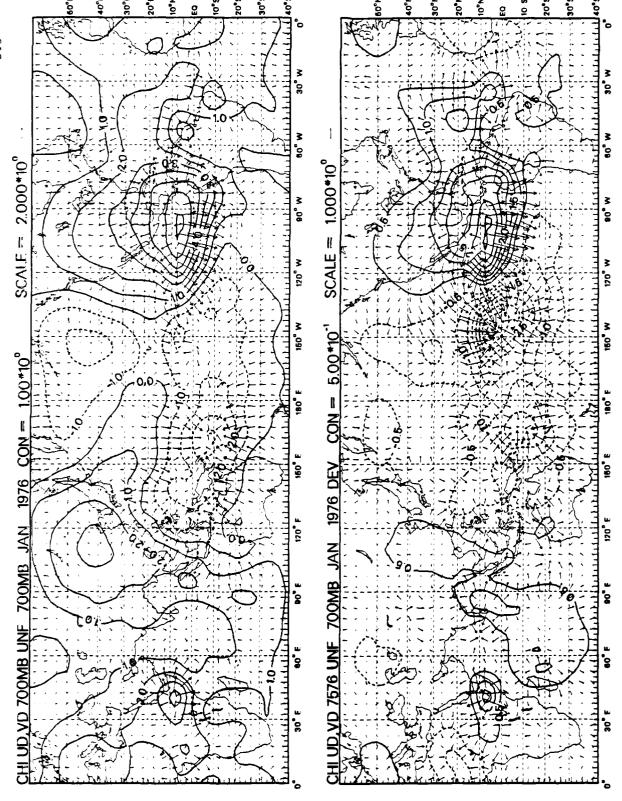


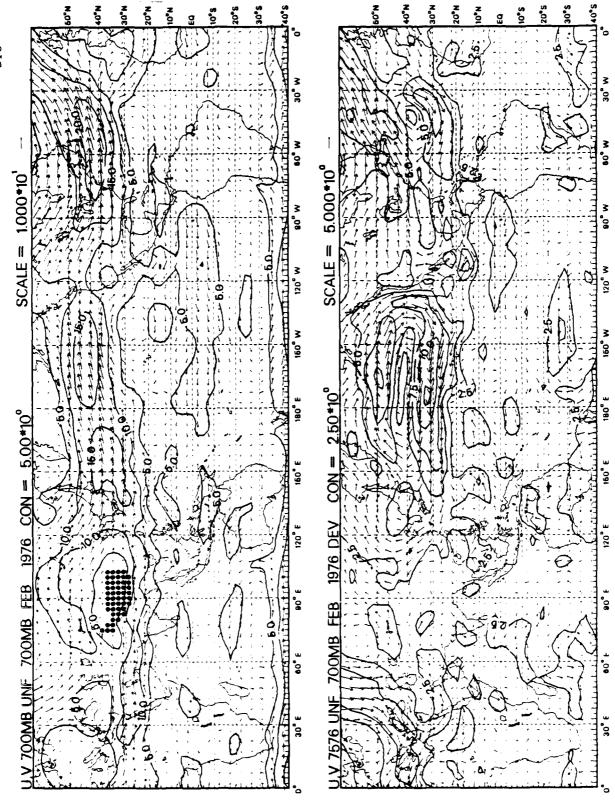


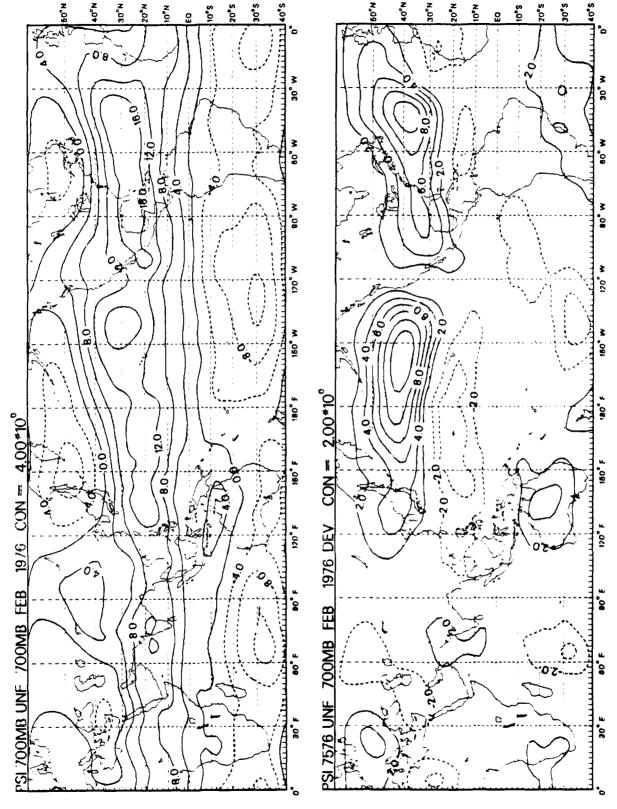


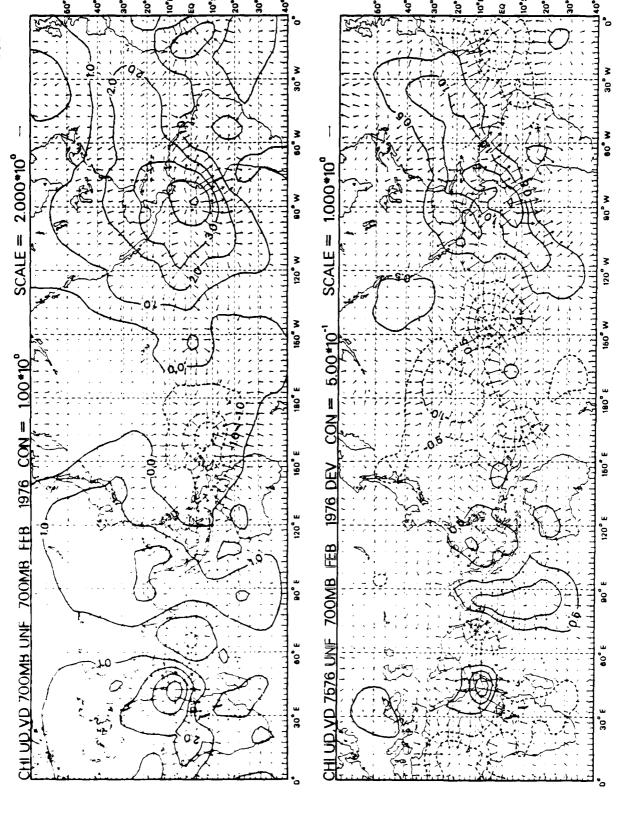


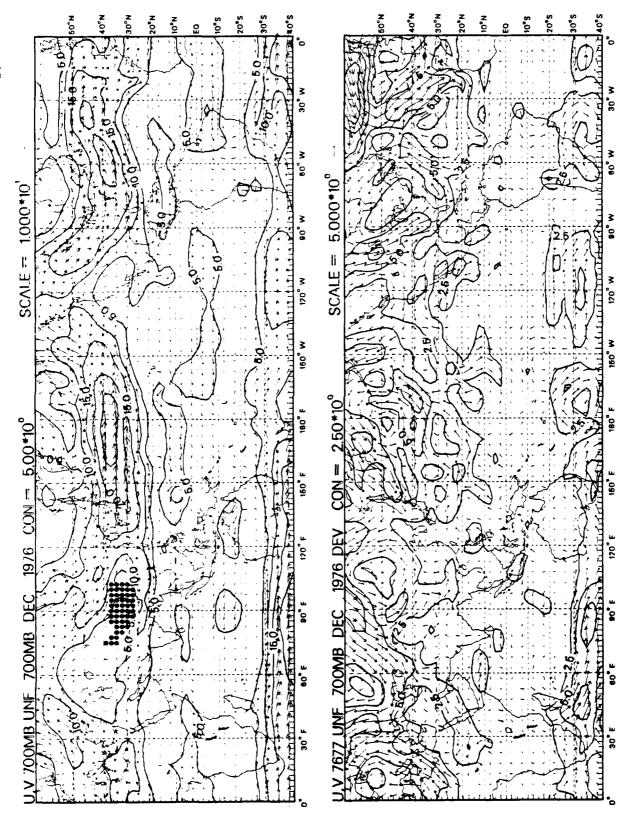


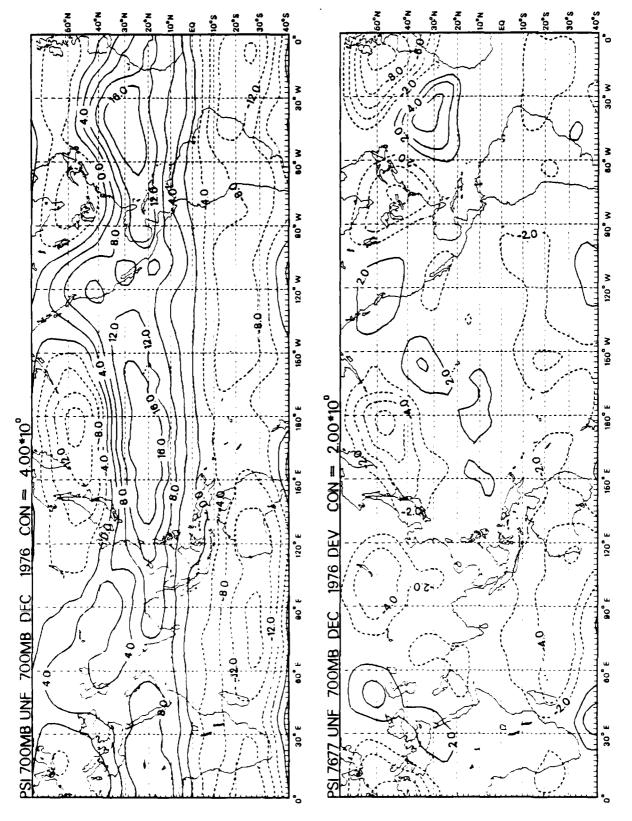


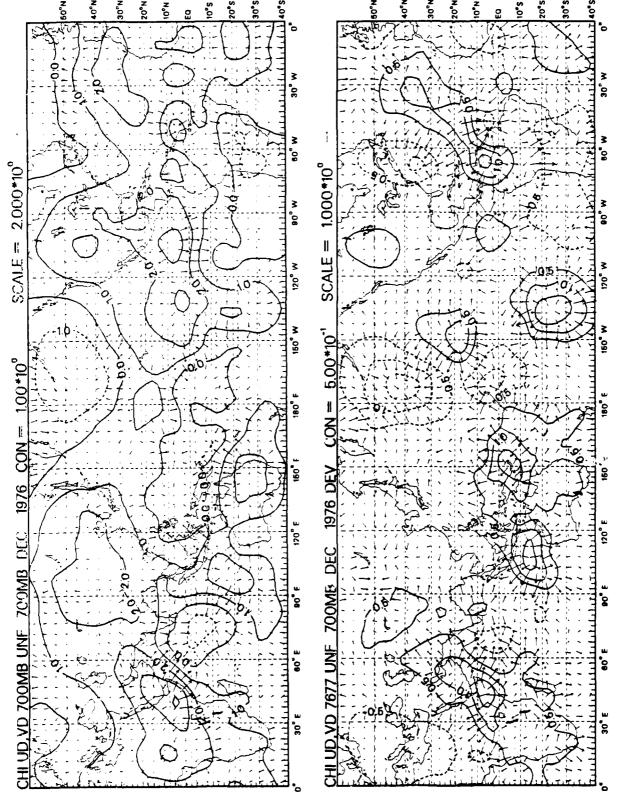


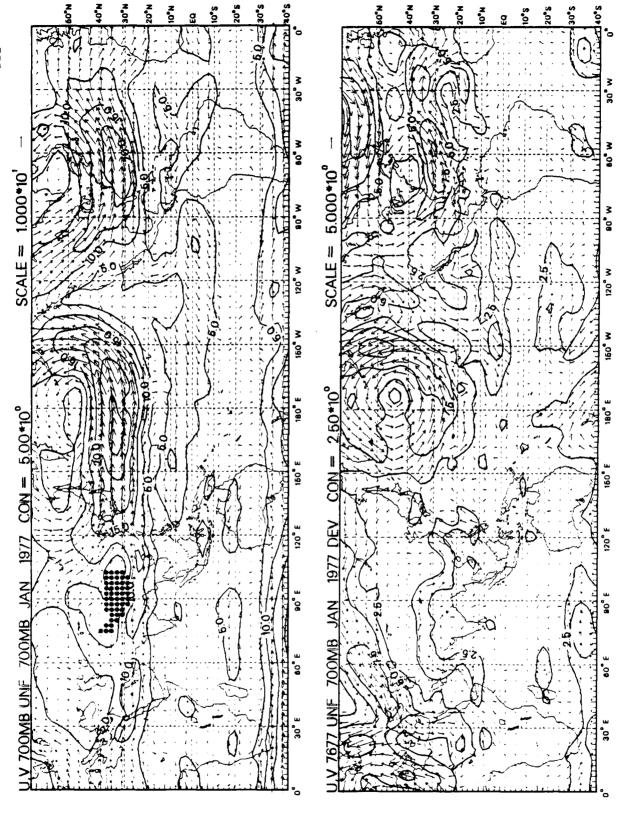


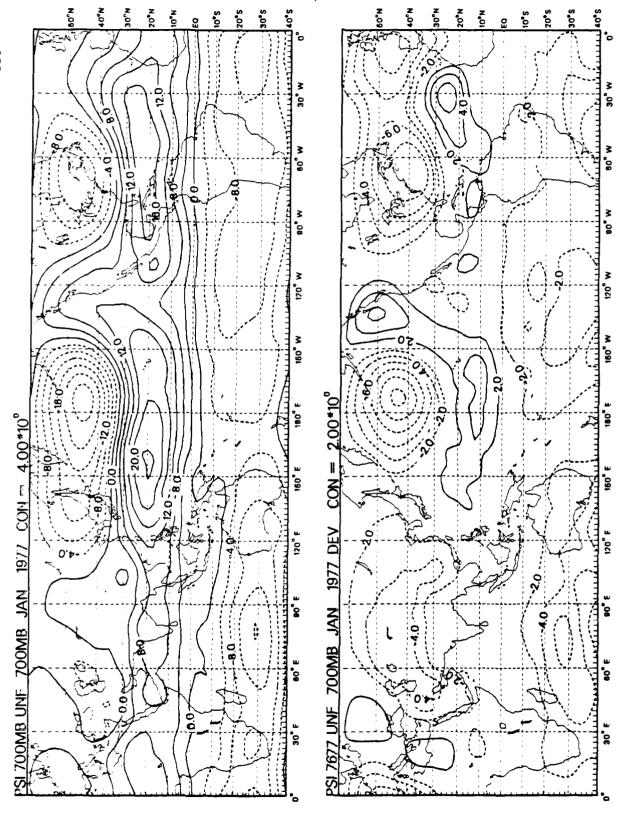


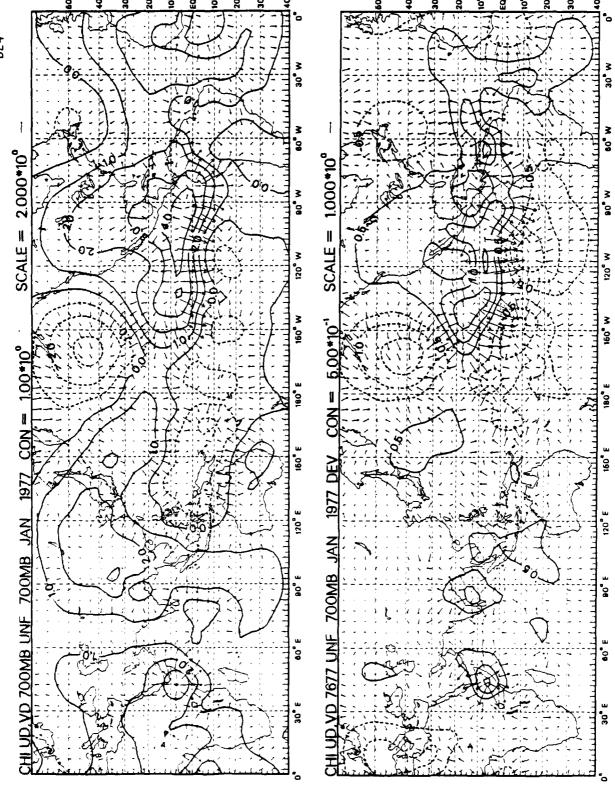


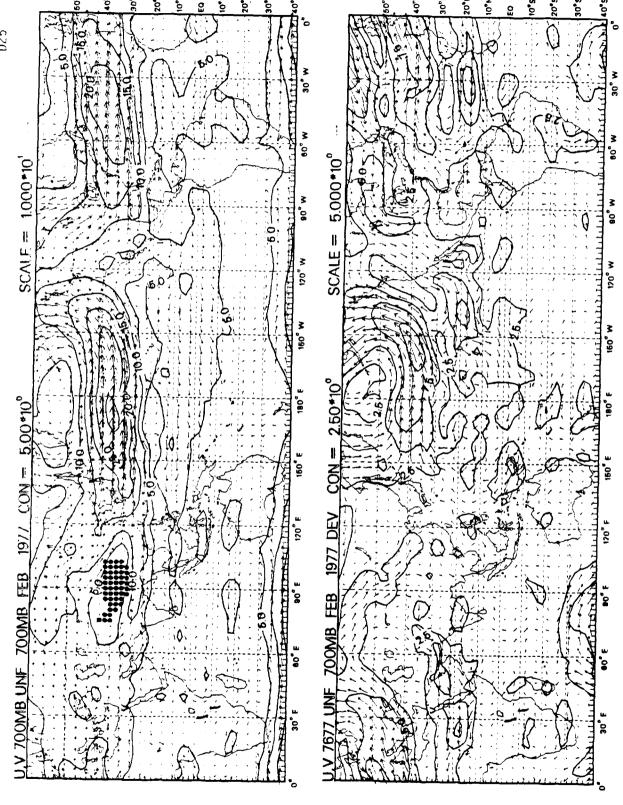


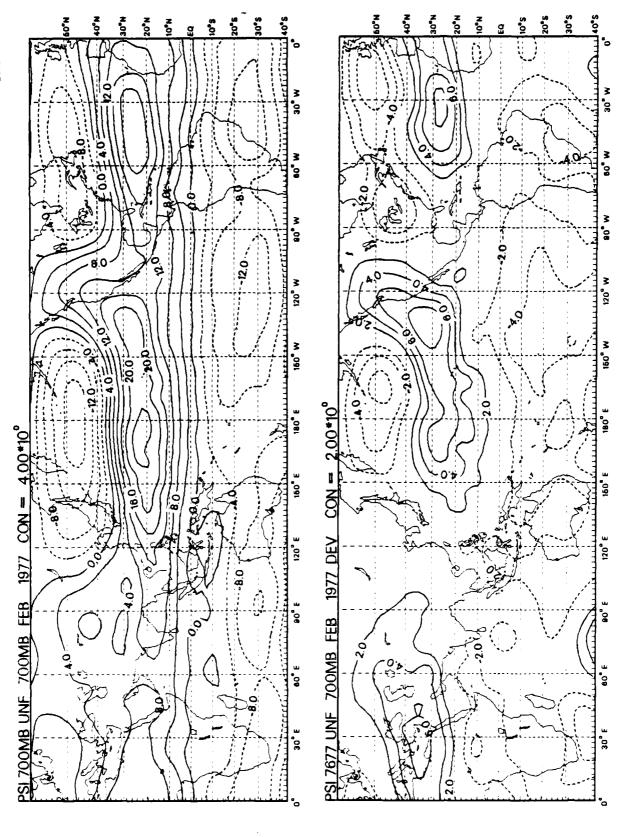


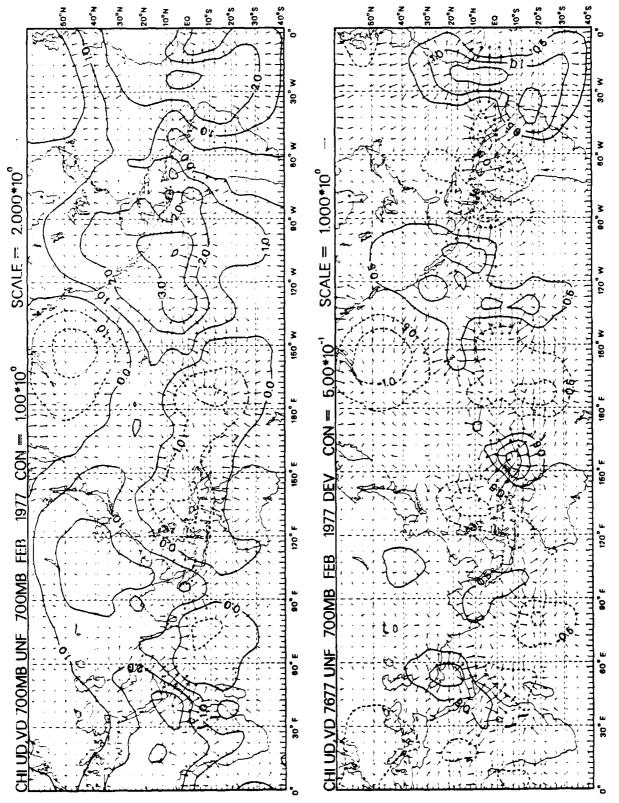


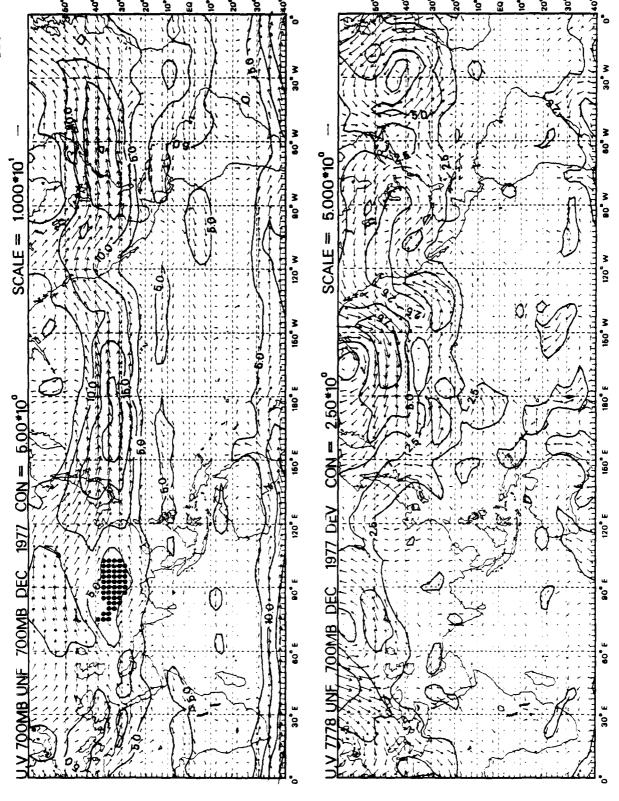


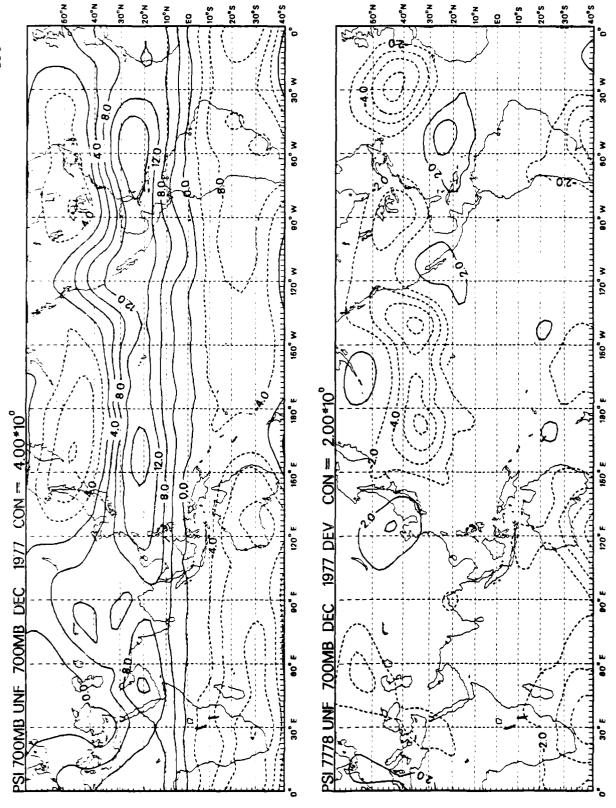


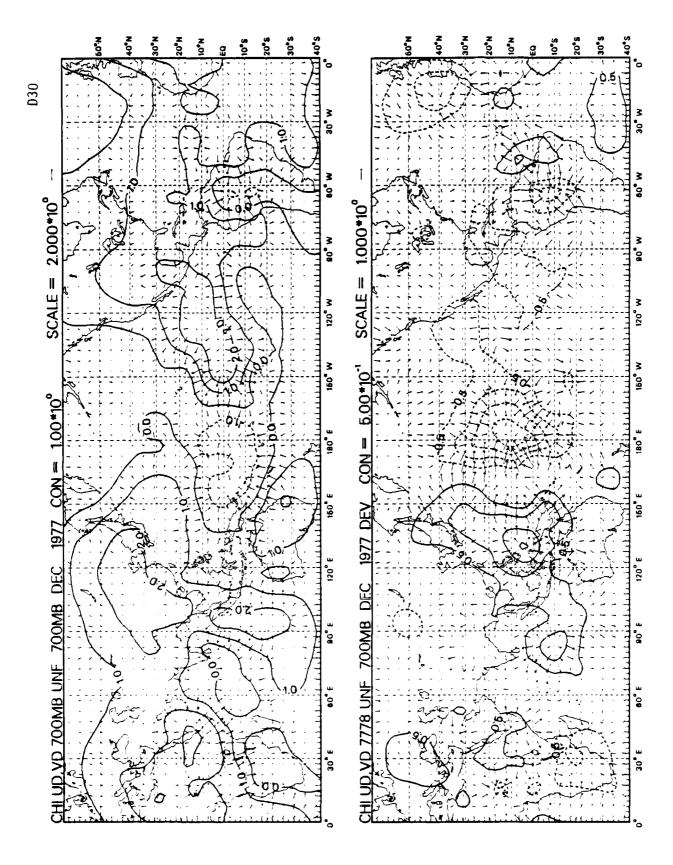


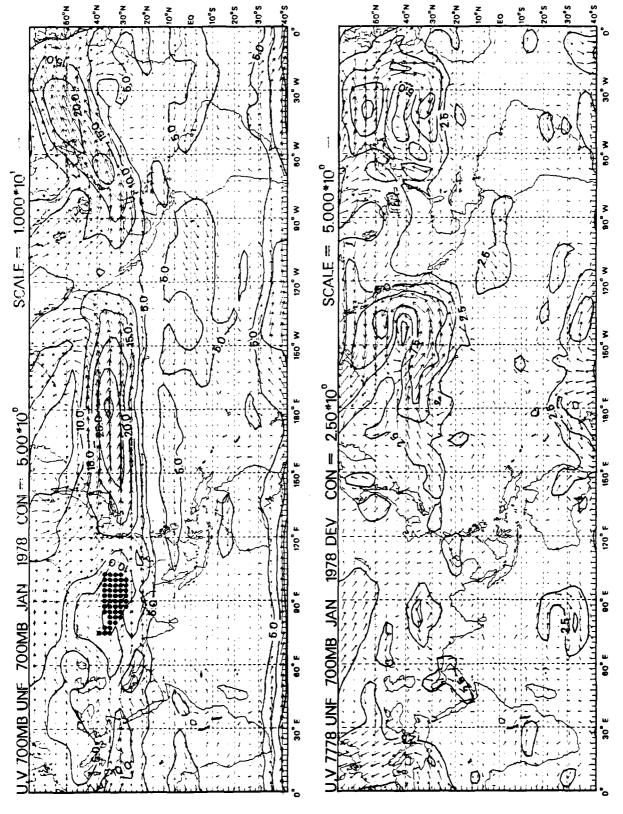


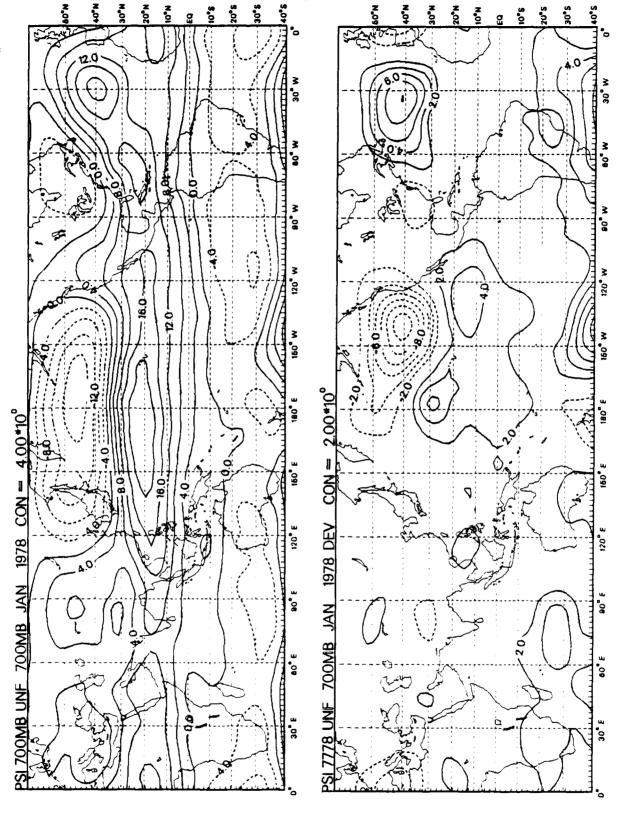




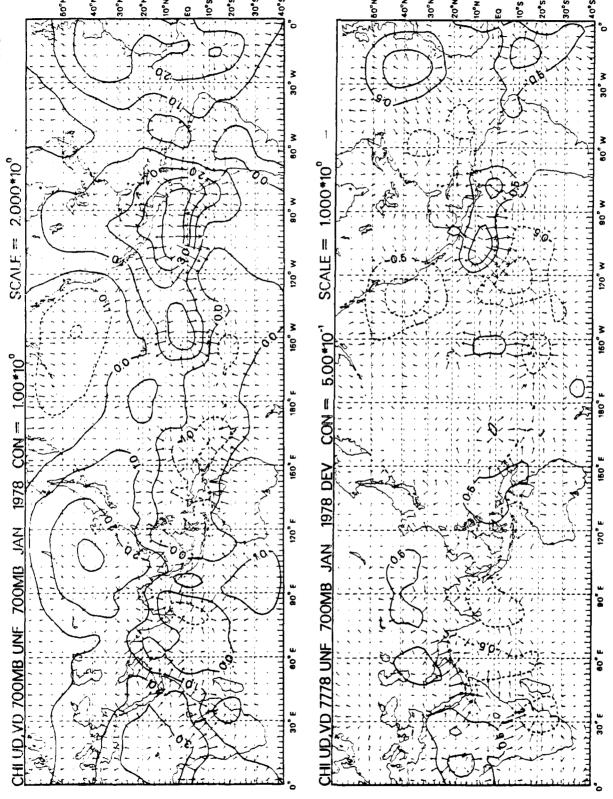




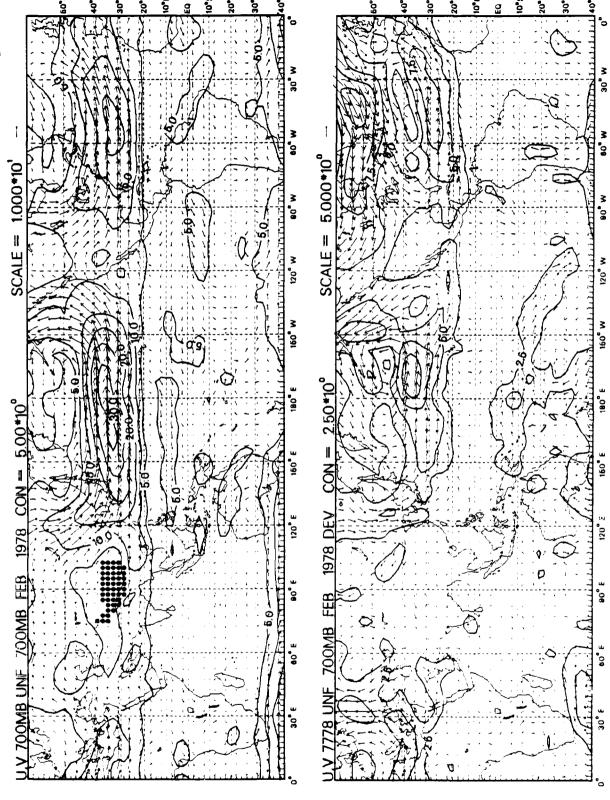












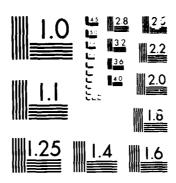
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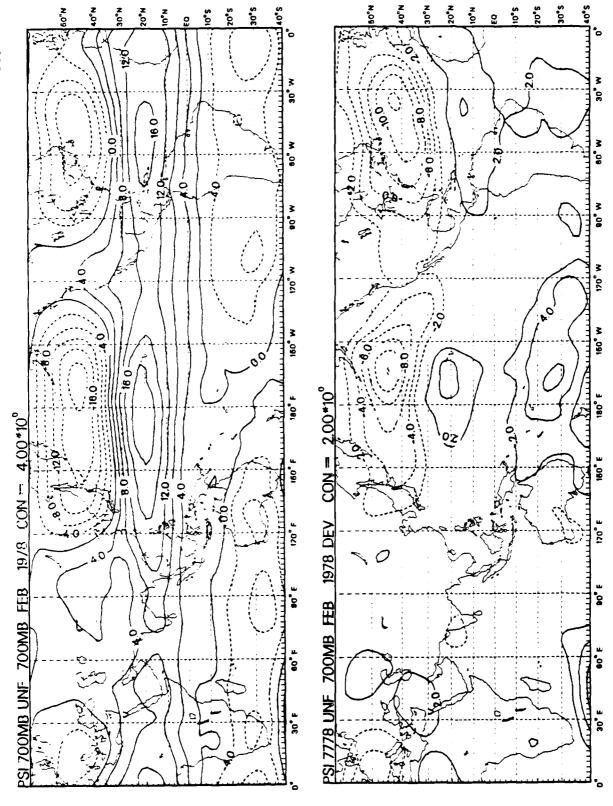
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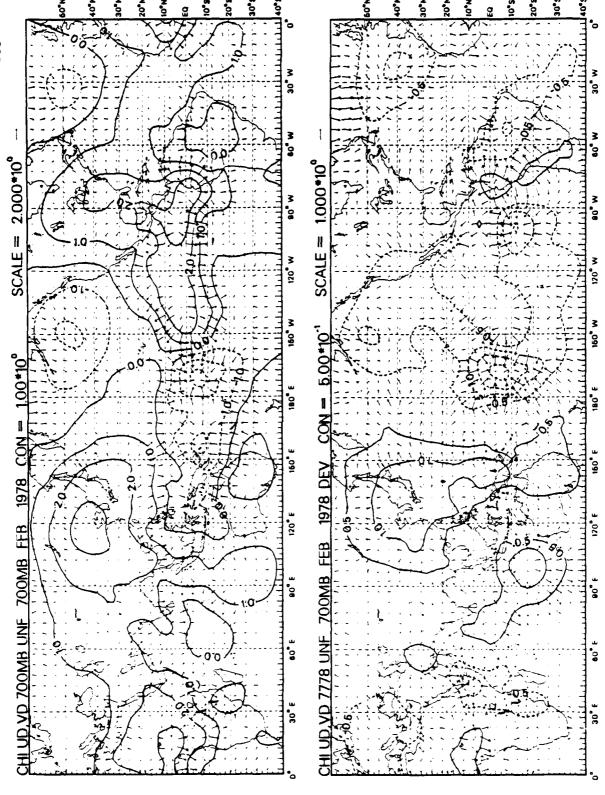


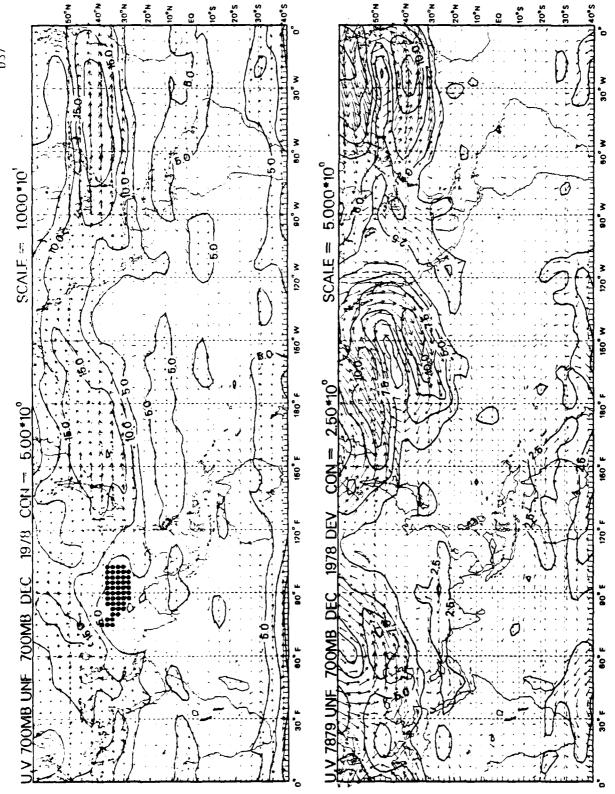
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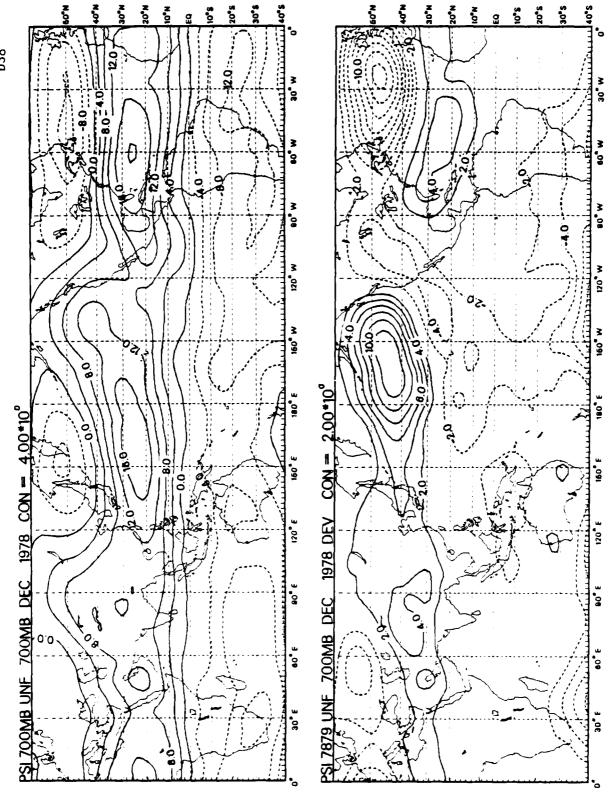


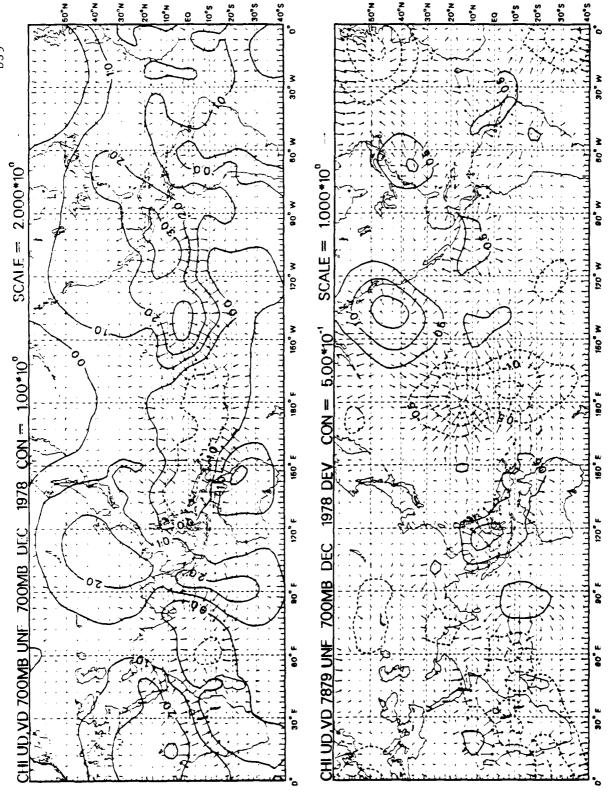




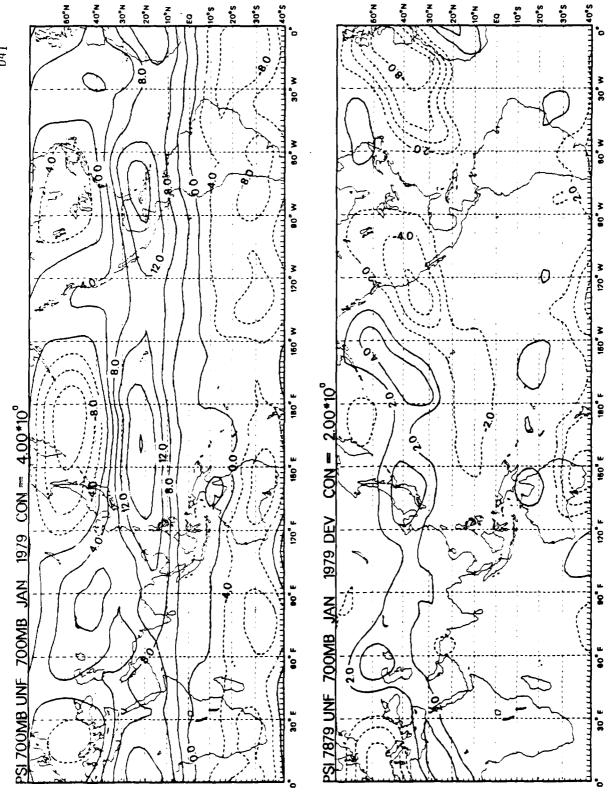


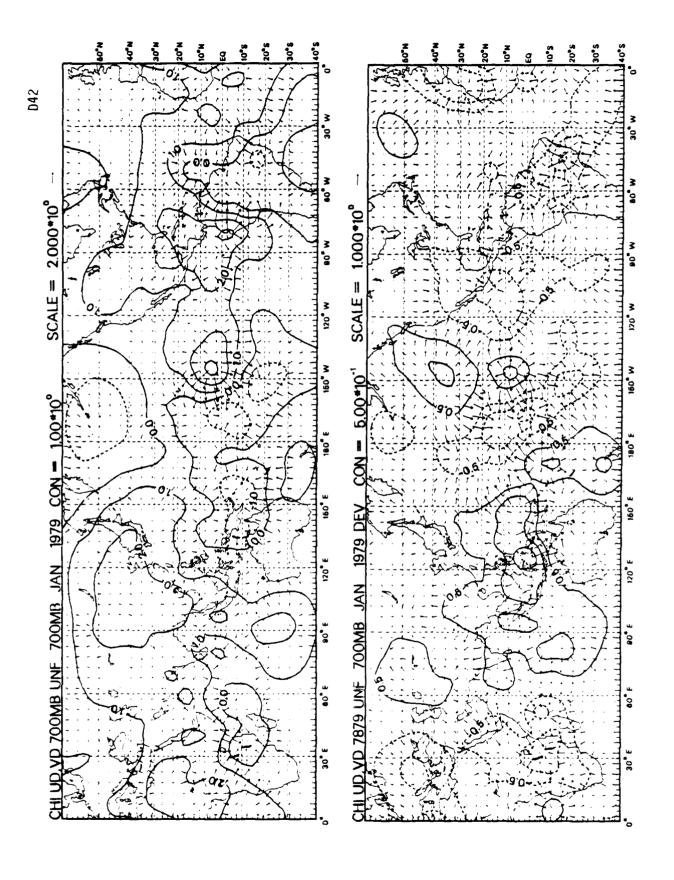


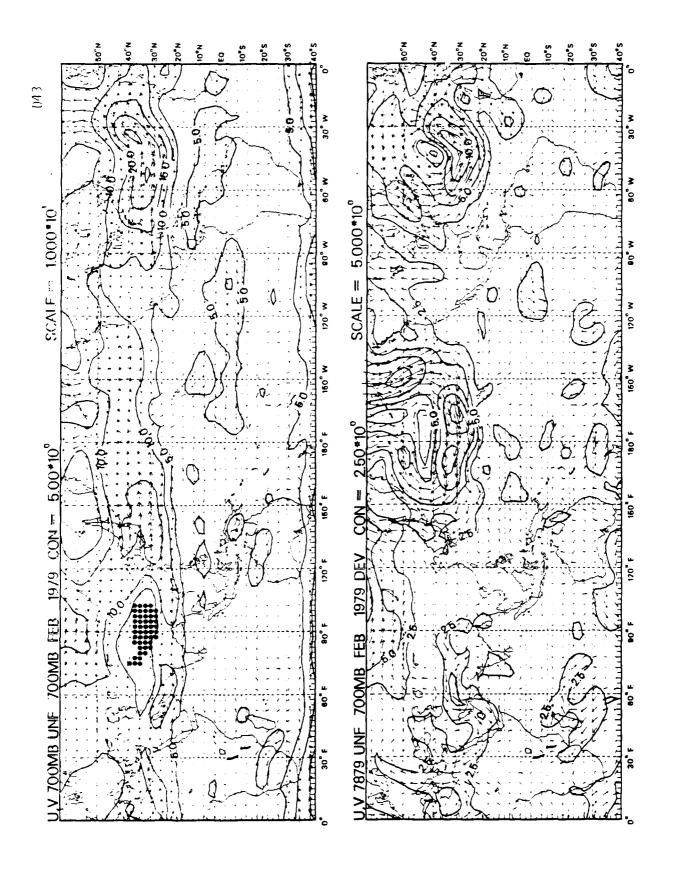


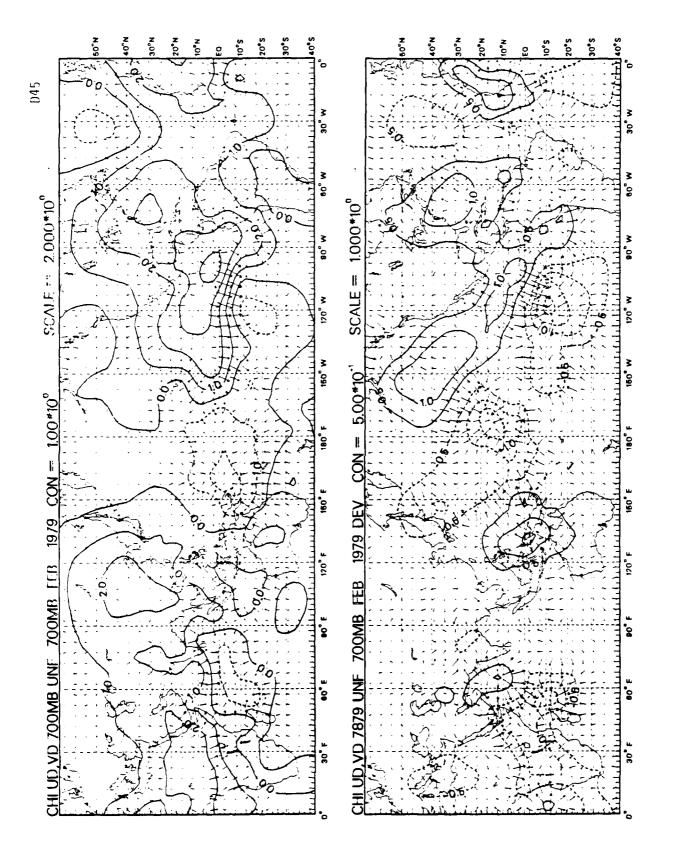


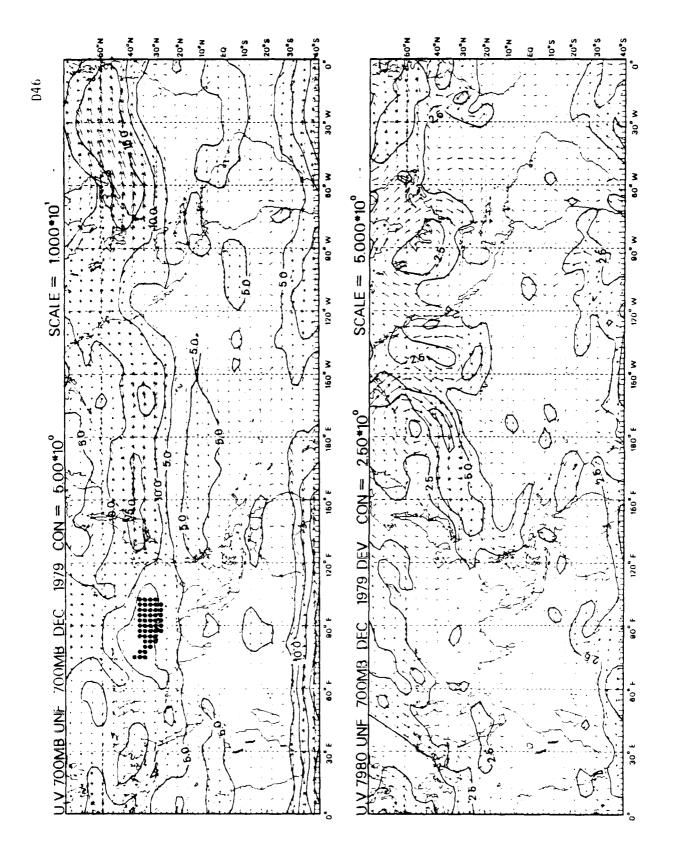


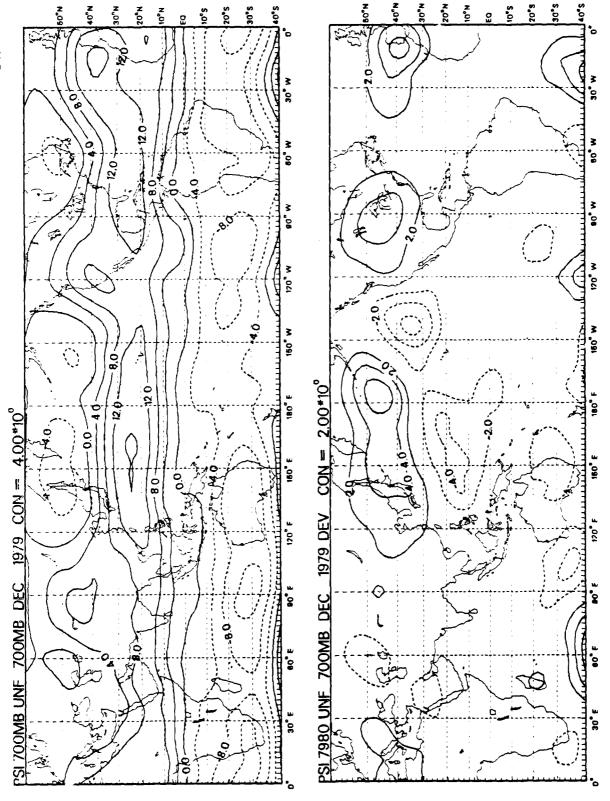


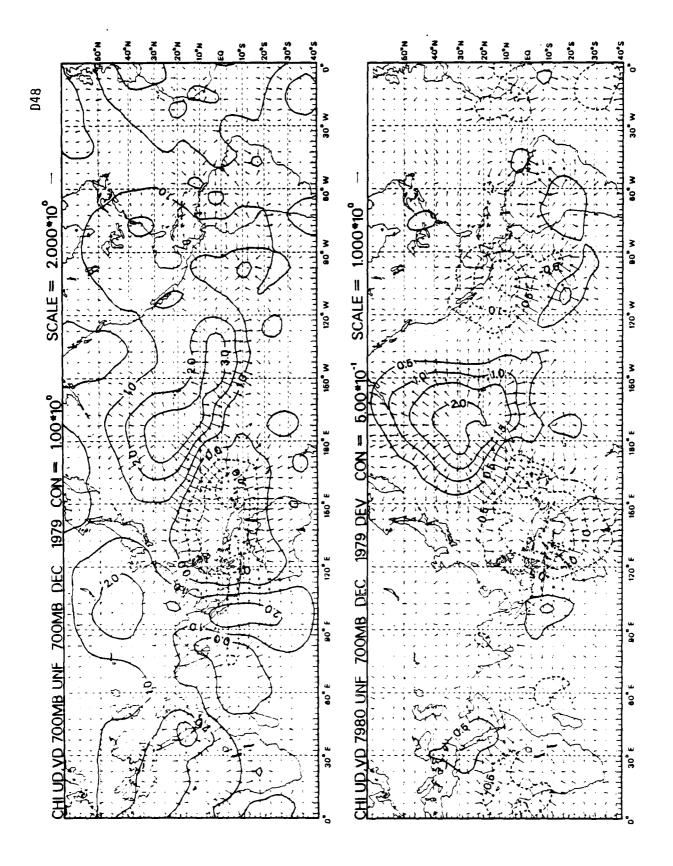


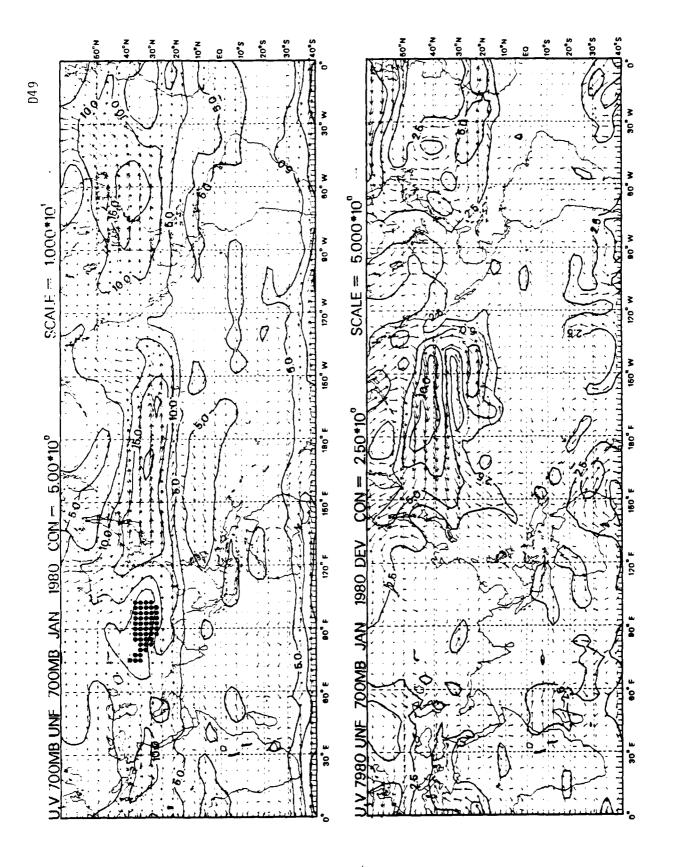


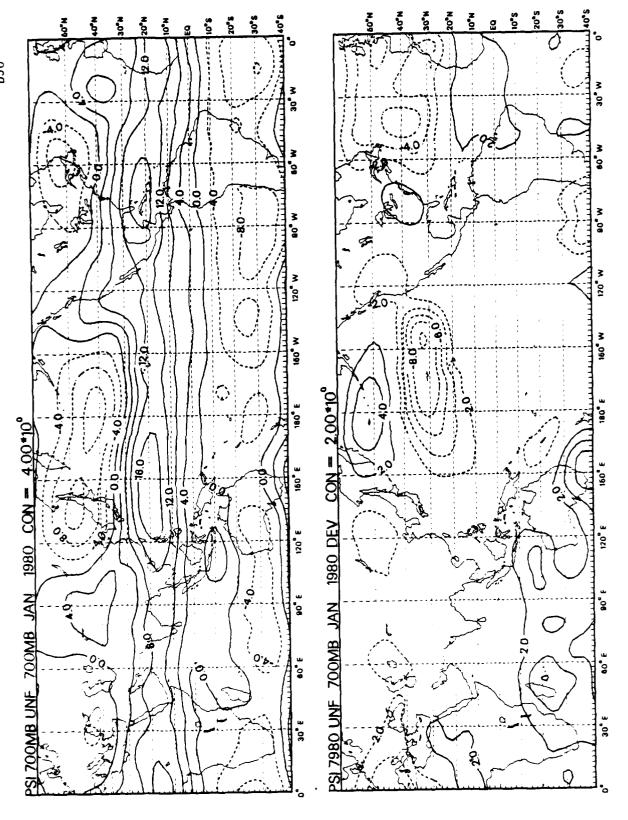




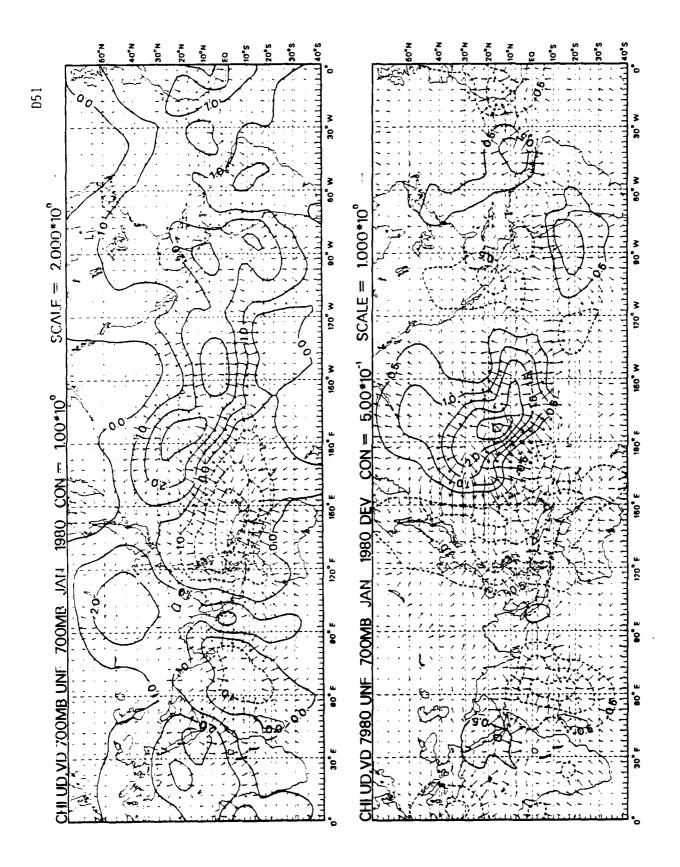


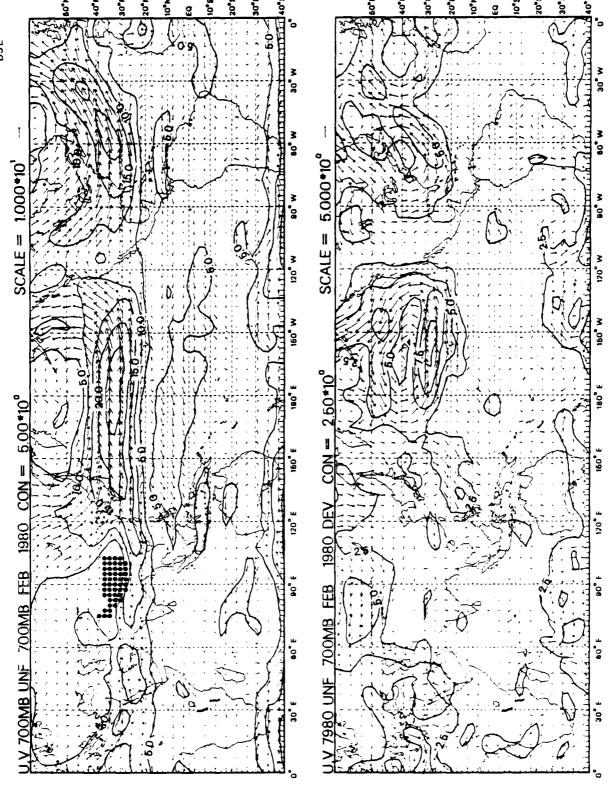


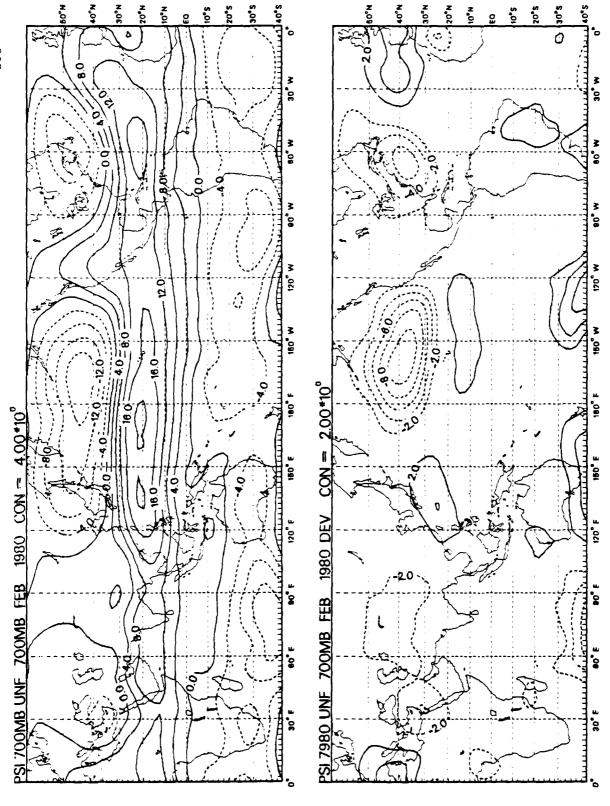


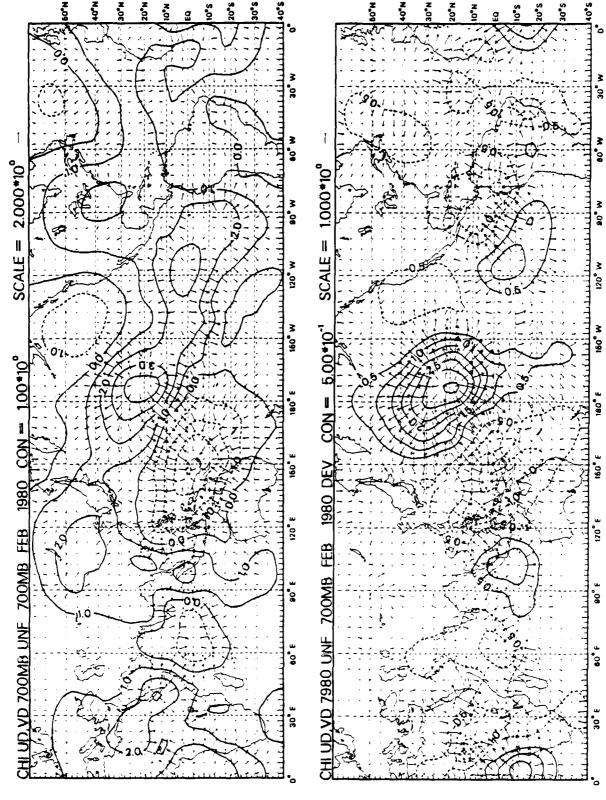


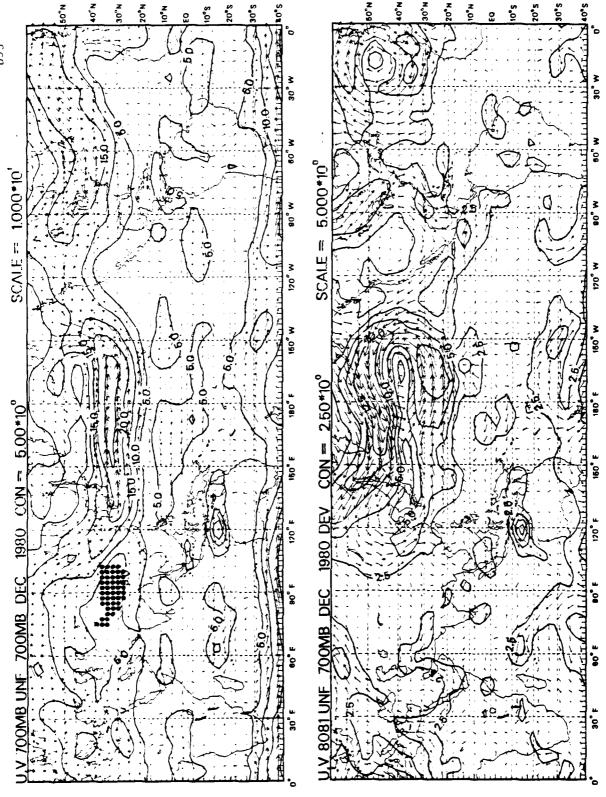
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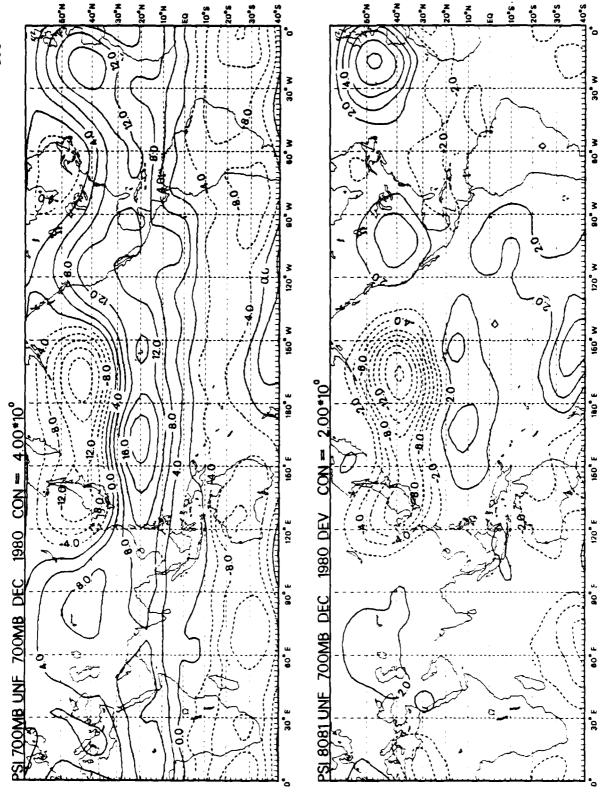


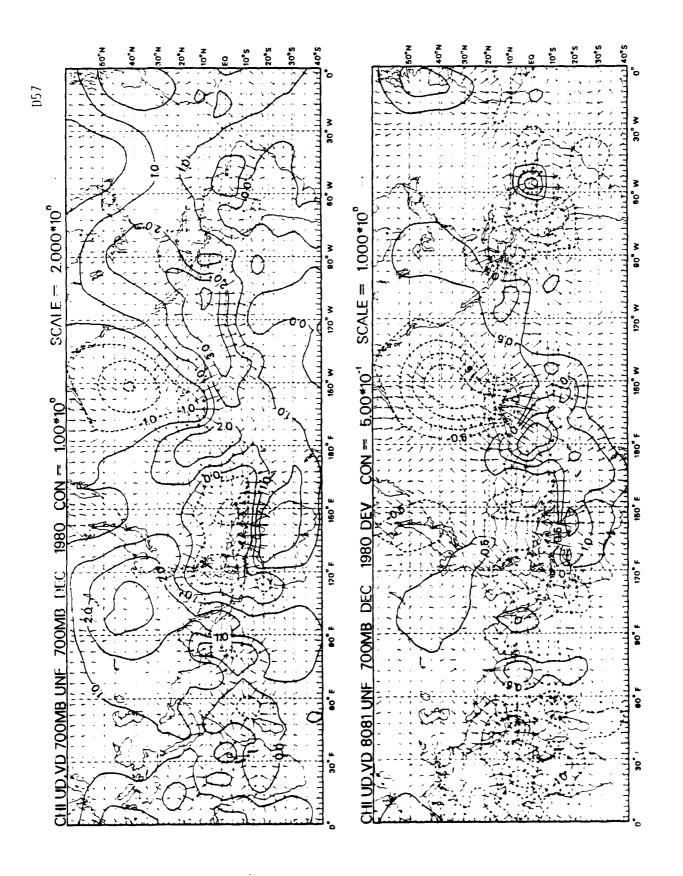


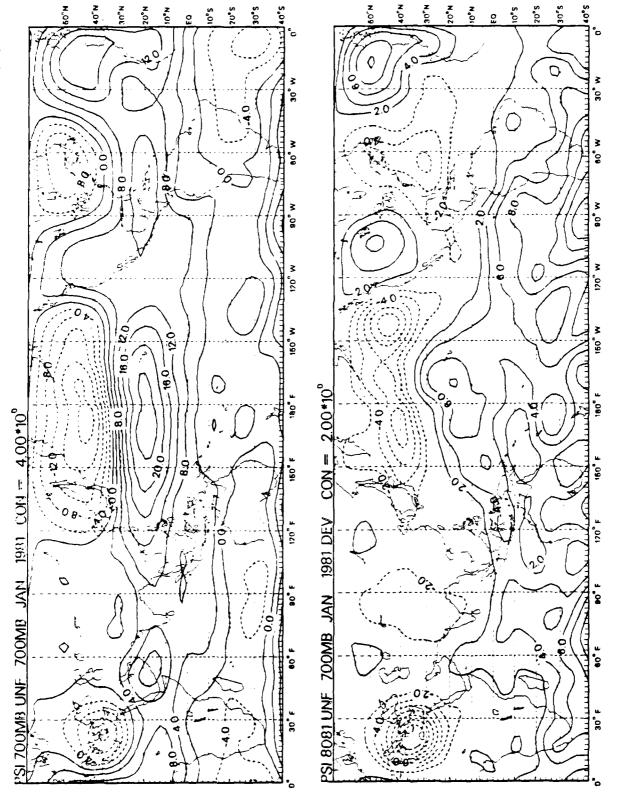


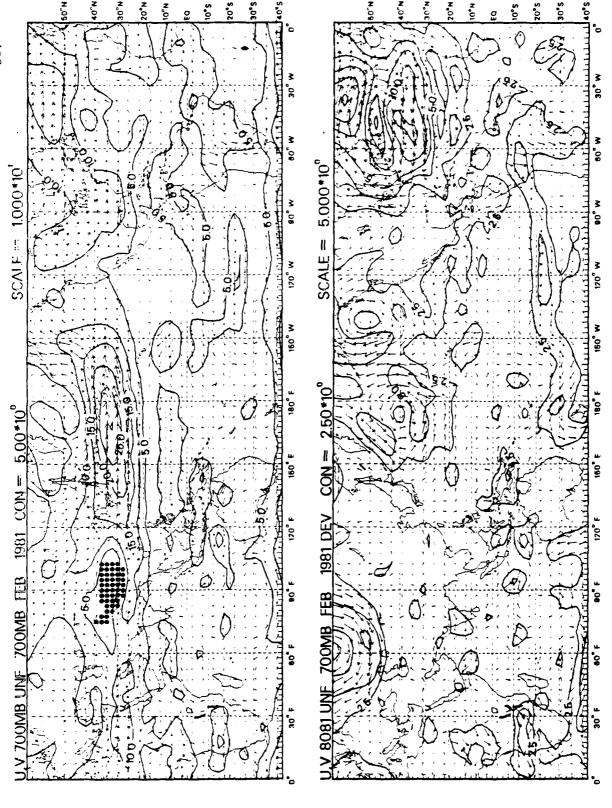






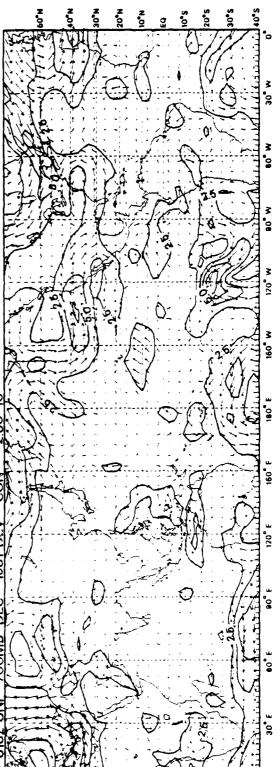


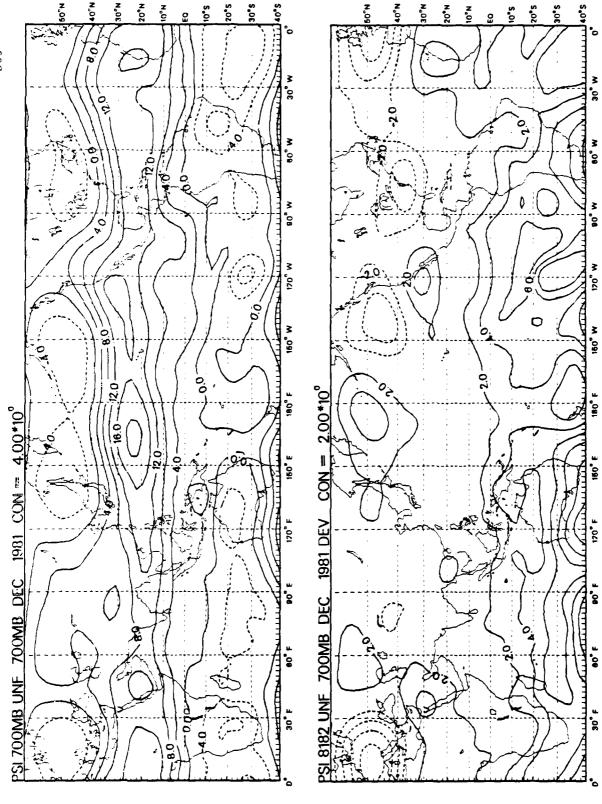


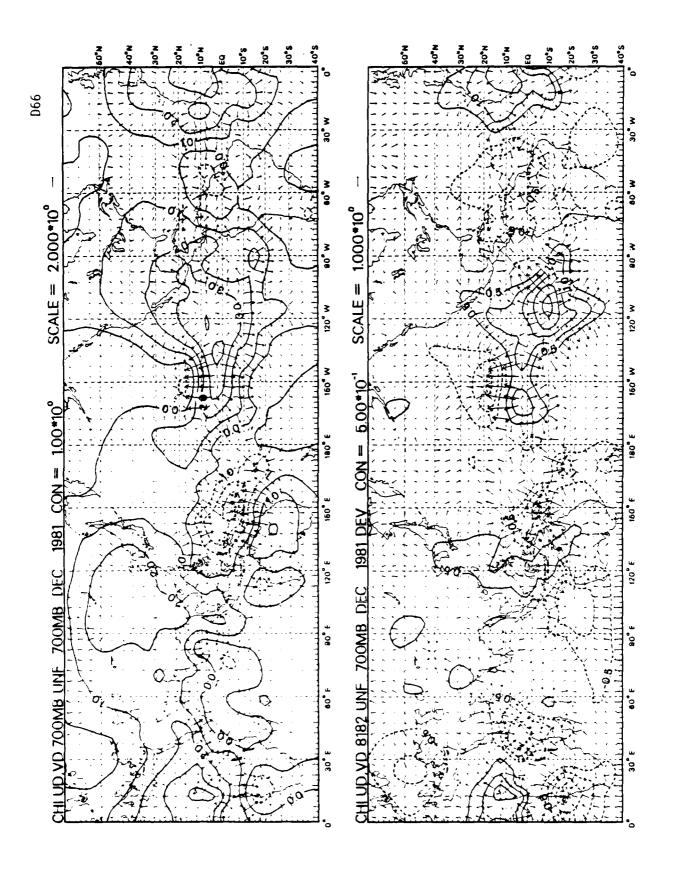


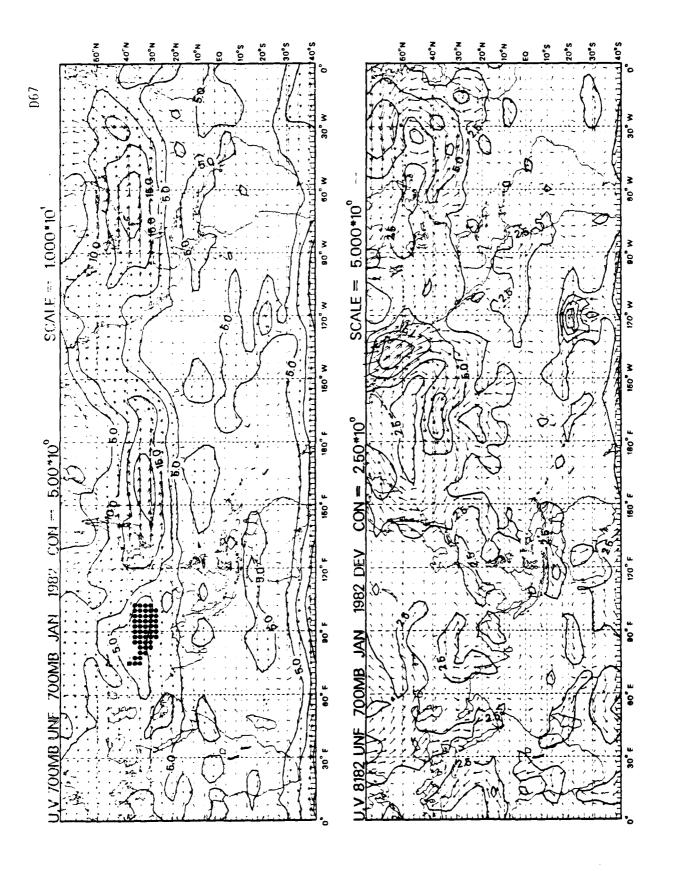
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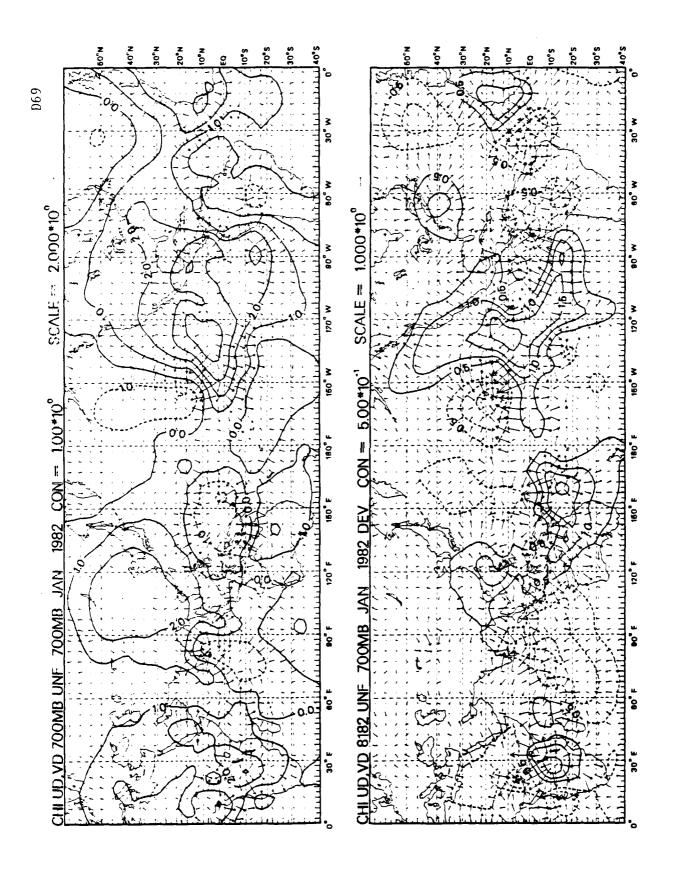
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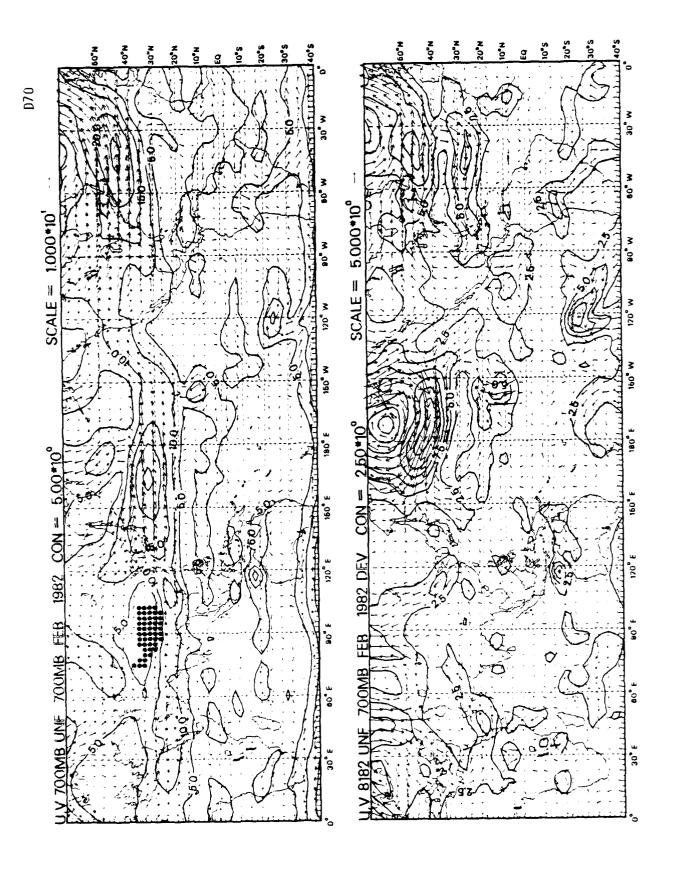


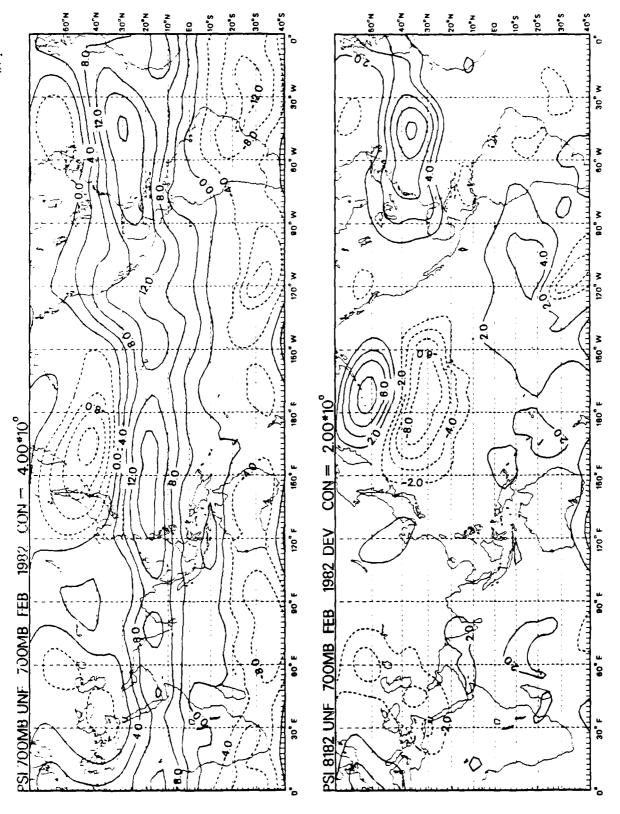


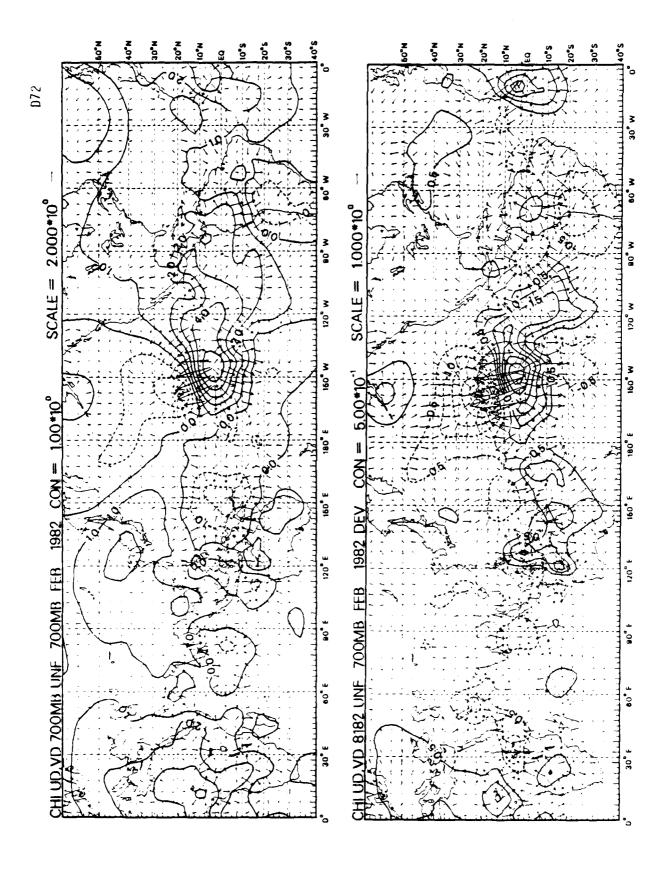


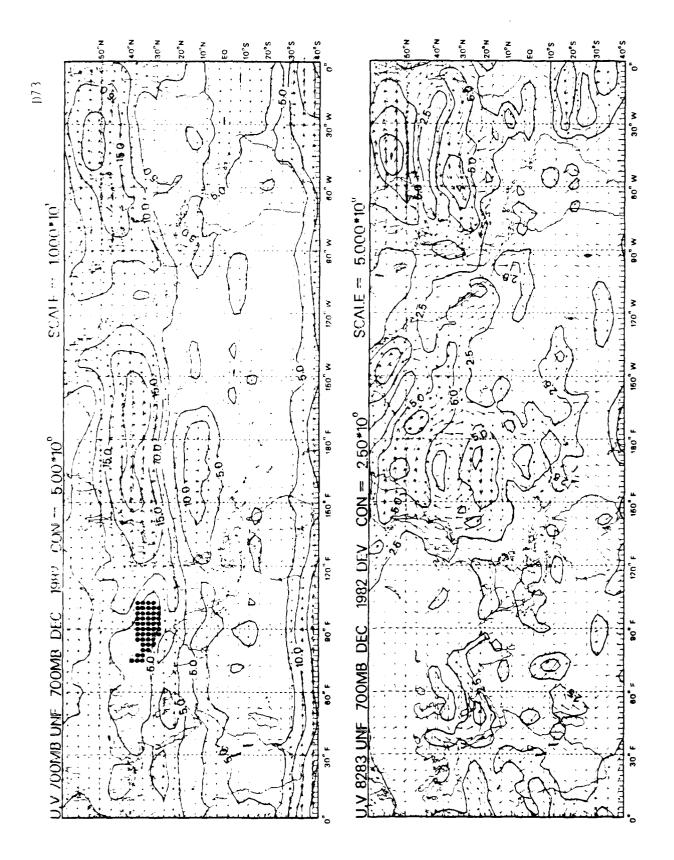


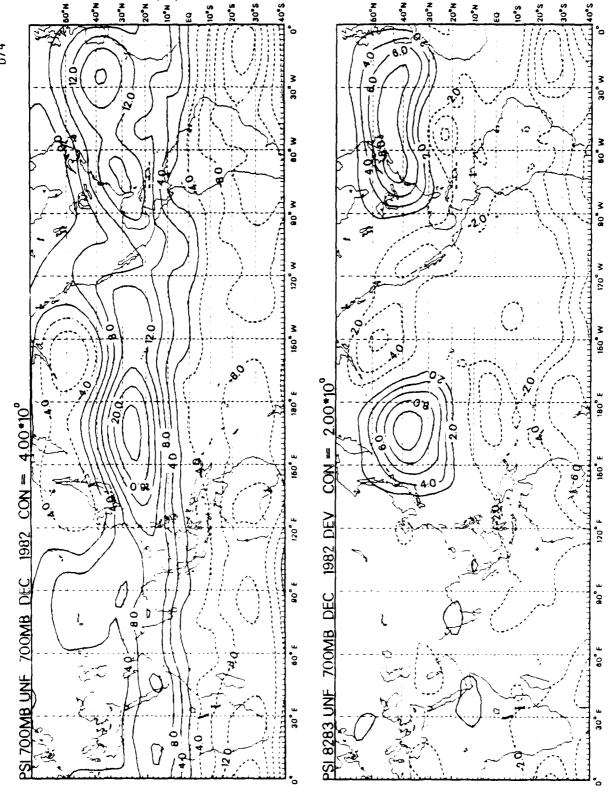


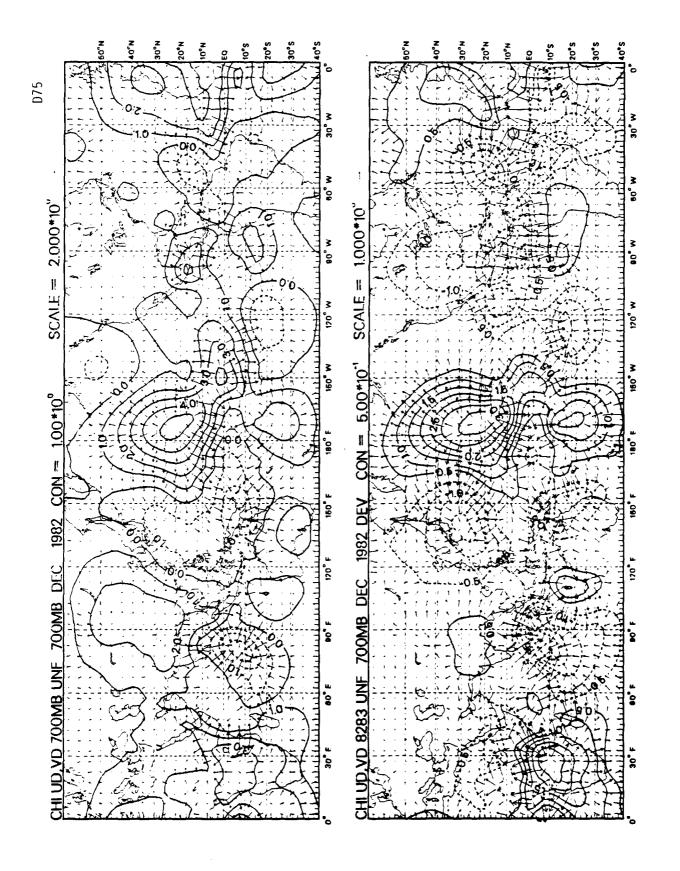


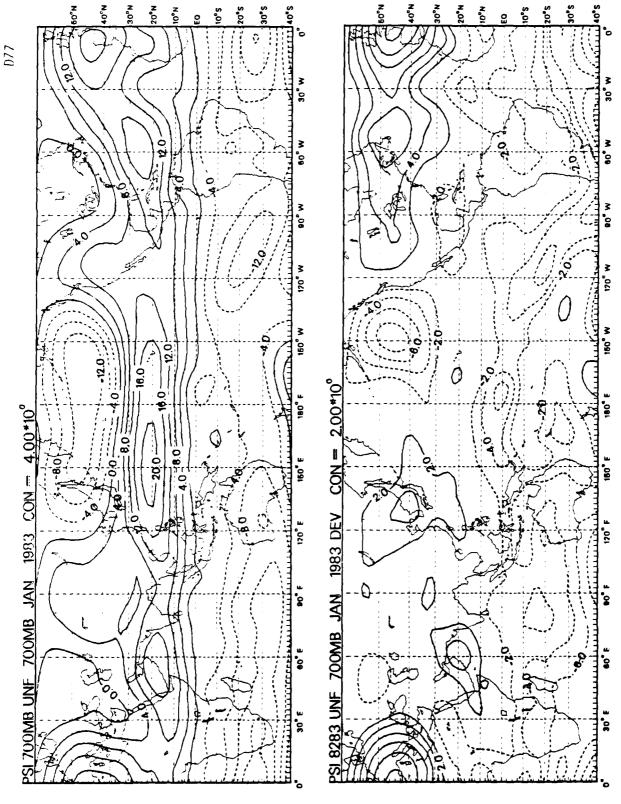


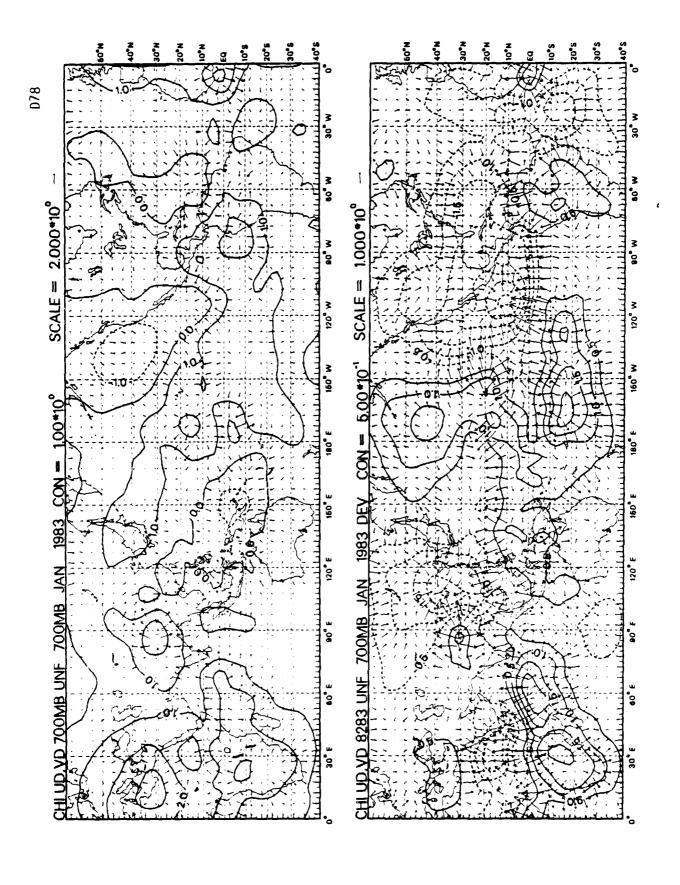


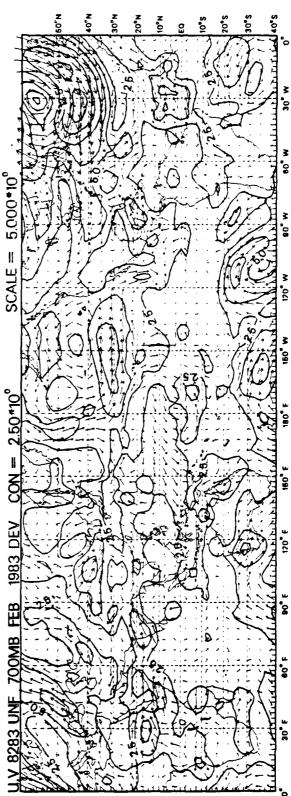


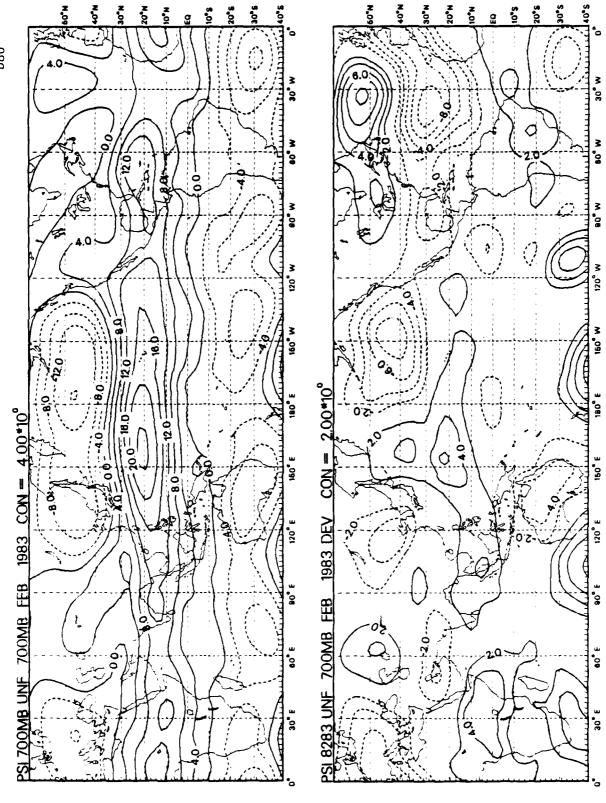


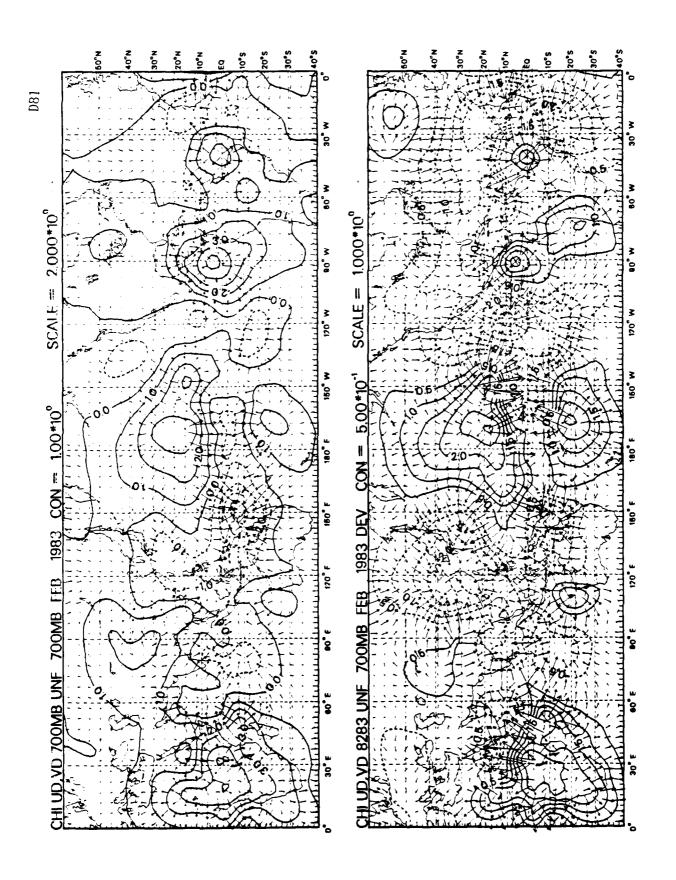












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